

JGR Atmospheres

RESEARCH ARTICLE

10.1029/2019JD031940

†Deceased 2015

Key Points:

- Downward Terrestrial Gamma-ray Flashes occur during strong initial breakdown pulses of negative cloud-to-ground and cloud lightning
- The initial breakdown pulses consist of streamer-based fast negative breakdown having transient sub-pulse conducting events or “sparks”
- The streamer to leader transition of negative stepping occurs during strong currents in the final stage of initial breakdown pulses

Supporting Information:

- Figure S1

Correspondence to:

J. Remington,
jremington@cosmic.utah.edu

Citation:

Belz, J. W., Krehbiel, P. R., Remington, J., Stanley, M. A., Abbasi, R. U., LeVon, R., et al. (2020). Observations of the origin of downward terrestrial gamma-ray flashes. *Journal of Geophysical Research: Atmospheres*, 125, e2019JD031940. <https://doi.org/10.1029/2019JD031940>

Received 5 NOV 2019

Accepted 21 SEP 2020

Accepted article online 5 OCT 2020

Observations of the Origin of Downward Terrestrial Gamma-Ray Flashes

J. W. Belz¹, P. R. Krehbiel² , J. Remington¹ , M. A. Stanley² , R. U. Abbasi³, R. LeVon¹, W. Rison², D. Rodeheffer², T. Abu-Zayyad¹, M. Allen¹, E. Barcikowski¹, D. R. Bergman¹, S. A. Blake¹, M. Byrne¹, R. Cady¹, B. G. Cheon⁴, M. Chikawa⁵, A. di Matteo^{6,7} , T. Fujii⁸, K. Fujita⁹, R. Fujiwara⁹, M. Fukushima^{10,11}, G. Furlich¹ , W. Hanlon¹ , M. Hayashi¹², Y. Hayashi⁹, N. Hayashida¹³, K. Hibino¹³, K. Honda¹⁴, D. Ikeda¹⁴ , T. Inadomi¹⁶, N. Inoue¹⁷, T. Ishii¹⁴, H. Ito¹⁸, D. Ivanov¹, H. Iwakura¹⁶, H. M. Jeong¹⁹, S. Jeong¹⁹, C. C. H. Jui¹, K. Kadota²⁰, F. Kakimoto²¹, O. Kalashev²², K. Kasahara²³ , S. Kasami²⁴, H. Kawai²⁵, S. Kawakami⁹, K. Kawata¹⁰, E. Kido¹⁰, H. B. Kim⁴, J. H. Kim¹, J. H. Kim⁹ , V. Kuzmin^{†22}, M. Kuznetsov^{6,22}, Y. J. Kwon²⁶, K. H. Lee¹⁹, B. Lubsandorzhev²², J. P. Lundquist¹, K. Machida¹⁴, H. Matsumiya⁹, J. N. Matthews¹, T. Matuyama⁹, R. Mayta⁹, M. Minamino⁹, K. Mukai¹⁴, I. Myers¹, S. Nagataki¹⁸, K. Nakai⁹, R. Nakamura¹⁶, T. Nakamura²⁷, Y. Nakamura¹⁶, T. Nonaka¹⁰ , H. Oda⁹, S. Ogio^{9,28}, M. Ohnishi¹⁰, H. Ohoka¹⁰, Y. Oku²⁴, T. Okuda²⁹, Y. Omura⁹, M. Ono¹⁸, A. Oshima³⁰, S. Ozawa²³, I. H. Park¹⁹, M. Potts¹ , M. S. Pshirkov^{22,31} , D. C. Rodriguez¹, G. Rubtsov²², D. Ryu³² , H. Sagawa¹⁰, R. Sahara⁹, K. Saito¹⁰, Y. Saito¹⁶, N. Sakaki¹⁰, T. Sako¹⁰ , N. Sakurai⁹, K. Sano¹⁶, T. Seki¹⁶, K. Sekino⁹, F. Shibata¹⁴, T. Shibata¹⁰, H. Shimodaira¹⁰, B. K. Shin⁹, H. S. Shin¹⁰, J. D. Smith¹, P. Sokolsky¹, N. Sone¹⁶, B. T. Stokes¹, T. A. Stroman¹, Y. Takagi⁹, Y. Takahashi⁹, M. Takeda¹⁰, R. Takeishi¹⁹, A. Taketa¹⁵, M. Takita¹⁰, Y. Tameda²⁴, K. Tanaka³³, M. Tanaka³⁴ , Y. Tanoue⁹, S. B. Thomas¹, G. B. Thomson¹, P. Tinyakov^{6,22}, I. Tkachev²², H. Tokuno²¹, T. Tomida¹⁶, S. Troitsky²², Y. Tsunesada^{9,28} , Y. Uchihori³⁵, S. Udo¹³, T. Uehama¹⁶, F. Urban³⁶, M. Wallace¹, T. Wong¹, M. Yamamoto¹⁶, H. Yamaoka³⁴, K. Yamazaki³⁰, K. Yashiro³⁷, M. Yosei²⁴, H. Yoshii³⁸, Y. Zhezher^{10,22} and Z. Zundel¹

¹Department of Physics and Astronomy, University of Utah, Salt Lake City, UT, USA, ²Langmuir Laboratory for Atmospheric Research, New Mexico Institute of Mining and Technology, Socorro NM, USA, ³Department of Physics, Loyola University Chicago, Chicago, IL, USA, ⁴Department of Physics and The Research Institute of Natural Science, Hanyang University, Seoul, Korea, ⁵Department of Physics, Kindai University, Osaka, Japan, ⁶Service de Physique Théorique, Université Libre de Bruxelles, Brussels, Belgium, ⁷INFN, sezione di Torino, Turin, Italy, ⁸The Hakubi Center for Advanced Research, Kyoto University, Kitashirakawa-Oiwakecho, Kyoto, Japan, ⁹Graduate School of Science, Osaka City University, Osaka, Japan, ¹⁰Institute for Cosmic Ray Research, University of Tokyo, Kashiwa, Japan, ¹¹Kavli Institute for the Physics and Mathematics of the Universe (WPI), University of Tokyo, Kashiwa, Japan, ¹²Information Engineering Graduate School of Science and Technology, Shinshu University, Nagano, Japan, ¹³Faculty of Engineering, Kanagawa University, Yokohama, Japan, ¹⁴Interdisciplinary Graduate School of Medicine and Engineering, University of Yamanashi, Kofu, Japan, ¹⁵Earthquake Research Institute, University of Tokyo, Tokyo, Japan, ¹⁶Academic Assembly School of Science and Technology Institute of Engineering, Shinshu University, Nagano, Japan, ¹⁷The Graduate School of Science and Engineering, Saitama University, Saitama, Japan, ¹⁸Astrophysical Big Bang Laboratory, RIKEN, Wako, Japan, ¹⁹Department of Physics, Sungkyunkwan University, Suwon, Korea, ²⁰Department of Physics, Tokyo City University, Tokyo, Japan, ²¹Graduate School of Science and Engineering, Tokyo Institute of Technology, Tokyo, Japan, ²²Institute for Nuclear Research of the Russian Academy of Sciences, Moscow, Russia, ²³Advanced Research Institute for Science and Engineering, Waseda University, Tokyo, Japan, ²⁴Department of Engineering Science, Faculty of Engineering, Osaka Electro-Communication University, Neyagawa-shi, Japan, ²⁵Department of Physics, Chiba University, Chiba, Japan, ²⁶Department of Physics, Yonsei University, Seoul, Korea, ²⁷Faculty of Science, Kochi University, Kochi, Japan, ²⁸Nambu Yoichiro Institute of Theoretical and Experimental Physics, Osaka City University, Osaka, Japan, ²⁹Department of Physical Sciences, Ritsumeikan University, Kusatsu, Japan, ³⁰Engineering Science Laboratory, Chubu University, Kasugai, Japan, ³¹Sternberg Astronomical Institute, Moscow M.V. Lomonosov State University, Moscow, Russia, ³²Department of Physics, Ulsan National Institute of Science and Technology, UNIST-gil, Ulsan, Korea, ³³Graduate School of Information Sciences, Hiroshima City University, Hiroshima, Japan, ³⁴Institute of Particle and Nuclear Studies, KEK, Tsukuba, Japan, ³⁵National Institute of Radiological Science, Chiba, Japan, ³⁶Institute of Physics, Czech Academy of Sciences, CEICO, Prague, Czech Republic, ³⁷Department of Physics, Tokyo University of Science, Noda, Japan, ³⁸Department of Physics, Ehime University, Matsuyama, Japan

Abstract In this paper we report the first close, high-resolution observations of downward-directed terrestrial gamma-ray flashes (TGFs) detected by the large-area Telescope Array cosmic ray observatory, obtained in conjunction with broadband VHF interferometer and fast electric field change measurements of the parent discharge. The results show that the TGFs occur during strong initial breakdown pulses (IBPs) in the first few milliseconds of negative cloud-to-ground and low-altitude intracloud flashes and that the IBPs are produced by a newly identified streamer-based discharge process called fast negative breakdown. The observations indicate the relativistic runaway electron avalanches (RREAs) responsible for producing the TGFs are initiated by embedded spark-like transient conducting events (TCEs) within the fast streamer system and potentially also by individual fast streamers themselves. The TCEs are inferred to be the cause of impulsive sub-pulses that are characteristic features of classic IBP sferics. Additional development of the avalanches would be facilitated by the enhanced electric field ahead of the advancing front of the fast negative breakdown. In addition to showing the nature of IBPs and their enigmatic sub-pulses, the observations also provide a possible explanation for the unsolved question of how the streamer to leader transition occurs during the initial negative breakdown, namely, as a result of strong currents flowing in the final stage of successive IBPs, extending backward through both the IBP itself and the negative streamer breakdown preceding the IBP.

1. Introduction

The interplay between lightning and high-energy particle physics was realized over two decades ago with the serendipitous observation of gamma radiation emanating from the Earth. The Burst and Transient Source Experiment (BATSE) instrument aboard NASA's Compton Gamma-Ray Observatory was designed to detect radiation from Gamma Ray Bursts (GRBs), deep-space events which are considered the most intense sources of electromagnetic radiation in the Universe. In 1994, BATSE unexpectedly recorded a series of brief, intense flashes of gamma rays, which appeared to originate at high altitudes (≥ 15 km above ground level) above thunderstorm regions (Carlson et al., 2007; Fishman et al., 1994). The terrestrial gamma-ray flashes (TGFs) lasted from hundreds of microseconds up to a millisecond or more, and their energy spectrum was consistent with bremsstrahlung emission from electrons with energies of several million electron volts (MeV) or greater.

Subsequent observations, now numbering in the thousands of events, aboard the Ramaty High Energy Solar Spectroscopic Imager (RHESSI) satellite (Gjesteland et al., 2012; Grefenstette et al., 2009), NASA's Fermi Gamma-ray Space Telescope (Briggs et al., 2013; Foley et al., 2014; Roberts et al., 2017), and the Astrorivelatore Gamma a Immagini Leggero (AGILE) satellite (Marisaldi et al., 2014) have shown that, instead of being produced at high altitude above storms, the TGFs originate at lower altitudes commensurate with being inside storms. In particular, it has been shown that the TGFs are produced at the altitudes of intracloud (IC) lightning flashes, during upward negative breakdown at the beginning of the flashes (Cummer et al., 2011, 2015; Lu et al., 2010; Lyu et al., 2016; Mailyan et al., 2016; Shao et al., 2010; Stanley et al., 2006). The early RHESSI observations were found to be associated with millisecond-duration initial breakdown activity that occurs in the beginning stages of IC flashes. However, a direct connection with the initial breakdown events was uncertain due to a 1–3 ms timing uncertainty in the RHESSI data (Lu et al., 2011).

In recent years, a small subset of TGFs has been associated with high-peak current (few hundred kiloampere) IC discharge events, called energetic in-cloud pulses (EIPs) (Lyu et al., 2015). EIPs are energetic versions of what are called preliminary or initial breakdown pulses (Marshall et al., 2013) that are characteristic features of the beginning stages of IC and negative cloud-to-ground (–CG) flashes. The EIP studies have utilized data from the Gamma-ray Burst Monitor (GBM) on Fermi (Briggs et al., 2010), which detects individual photons with microsecond timing accuracy, allowing more accurate correlation with ground-based low frequency (LF) radio atmospheric or “sferic” observations. Although EIPs are infrequent and the number of documented cases is small (a dozen or so), TGFs have been detected for 100% of EIPs that occurred within view of the Fermi satellite and within range of ground-based sferic sensors. As a result of this predictability, EIPs are considered to be high-probability producers of at least a class of TGF-generating lightning events (Cummer et al., 2017; Lyu et al., 2016, 2018). However, the detailed discharge processes that produce EIPs

have not been understood, due to the lack of measurements of the parent flashes with ground-based instrumentation (such observations of a close EIP by Tilles, 2020), reported while this paper was in review, provides the first detailed information on the discharge processes and storm environment that led to its occurrence, as discussed later).

As satellite-based observations of upward TGFs have accumulated, the question has been whether lightning produces downward TGFs that could be detected on the ground below or near thunderstorms. In particular, negative-polarity cloud-to-ground (–CG) discharges begin with downward negative breakdown that would be expected to produce TGFs directed earthward. Until recently, only a few TGFs had been detected at ground level in association with overhead lightning. Instead of being produced in the early stages of natural lightning, however, the gamma rays occurred either during the upward ascent of artificial trailing-wire, rocket-triggered lightning discharges (Dwyer, 2004; Hare et al., 2016) or at a later time in natural flashes, following high-current return strokes of –CG discharges (Dwyer et al., 2012; Ringuette et al., 2013; Tran et al., 2015). Also, a particularly strong downward TGF was recently reported during a winter thunderstorm by Wada et al. (2019) at the time of lightning discharge in the storm that appeared to be produced at low altitude (≈ 400 m) above ground. Otherwise, significant impediments to detecting downward TGFs have been (a) the increasingly strong attenuation of gamma radiation at low altitudes in the atmosphere, and (b) the ground-based detectors being either too far below and/or not widespread enough to detect the forward-beamed radiation. Both issues have been addressed with observations from the large-area (700 km²) Telescope Array Surface Detector (TASD) cosmic ray facility in central Utah.

In data collected between 2008 and 2013 there were 10 occasions in which the TASD was triggered by multiple bursts of energetic particles—not arising from cosmic rays. The events occurred within a millisecond of being detected by the U.S. National Lightning Detection Network (NLDN) (Abbasi et al., 2017), which identified them as being produced during –CG flashes. Follow-up observations with the TASD by the authors of the present study, obtained between 2014 and 2016 in coordination with a 3-D lightning mapping array (LMA) and a lightning electric field change sensor, detected 10 additional events, each consisting of three to five lightning-initiated bursts (Abbasi et al., 2018). The bursts were typically ≈ 10 μ s or less in duration and occurred over several hundred μ s time intervals during the first millisecond of downward negative breakdown at the beginning of –CG flashes. Scintillator responses and simulation studies showed that the bursts primarily resulted from gamma radiation and collectively comprised low-fluence TGFs. The LMA observations showed the bursts coincided with impulsive in-cloud VHF radiation events during energetic downward negative breakdown, 3–4 km above ground level. Although the TASD and LMA observations had sub-microsecond time resolution, the electric field change measurements recorded only the relatively slow electrostatic field change, with insufficient bandwidth to detect the faster electric field changes of the initial breakdown activity.

Here we report observations of downward TGFs produced by four additional flashes (three –CGs and one low-altitude IC flash) obtained in 2018 during continued studies with the Telescope Array. For this study, the TASD and LMA observations were augmented with crucially important, high-resolution VHF interferometric and fast electric field change measurements of the parent lightning discharges, obtained in relatively close proximity (16–24 km) to the TGFs. Coupled with sub-microsecond TGF measurements at TASD stations immediately below and near the flashes, the observations document the TGF occurrence with a high degree of temporal and spatial resolution not available before now. In each of the four flashes, the TGFs show a clear correspondence with downward negative breakdown during strong initial breakdown pulse (IBP) events in the first millisecond or so of the flashes. The negative breakdown progresses at a fast average speed ($\approx 1\text{--}3 \times 10^7$ m/s), indicative of a newly recognized type of discharge process called fast negative breakdown (FNB) (Tilles et al., 2019). Such breakdown is the negative analog of fast positive breakdown found in an earlier study to be the cause of high-power discharges called narrow bipolar events (NBEs) (Rison et al., 2016).

For both polarities, the breakdown is produced by a propagating system of streamers that substantially enhance (up to 50% or more) the electric field ahead of the streamers' advancing front (Attanasio et al., 2019). For the negative polarity version, electron avalanches produced within the streamer system would propagate through and ahead of the advancing front, producing downward-directed gamma radiation. Detailed analysis of the observations indicates that the TGFs are often initiated at the time of characteristic “sub-pulses” that occur during large-amplitude, “classic” sferics. From this, we infer that the sub-pulses are produced by

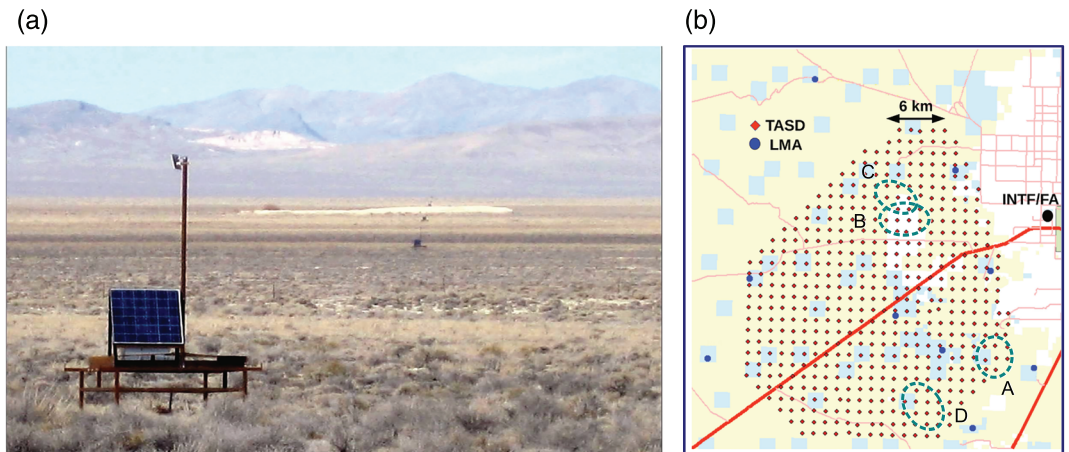


Figure 1. Telescope Array Surface Detector. (a) View of a close and distant surface detector stations on the desert plain west of Delta, Utah. Each detector unit consists of two 3 m^2 by 1.2 cm thick scintillator planes separated by a 0.1 cm steel sheet (Abu-Zayyad et al., 2013). Photo by M. Fukushima. (b) Map of the TASD stations, showing the locations of TGFs A–D (dashed ellipses). A total of 512 surface detectors have been deployed over a 700 km^2 area on a 1.2 km grid since 2008. A nine-station 3-D lightning mapping array (LMA) has been operated at the TASD since 2013 (blue dots). In July 2018, a VHF interferometer (INTF) and fast electric field spheric sensor (FA) were deployed 6 km east of the TASD, only a few days prior to observing the TGFs reported here.

transient spark-like discharges embedded within the negative streamer system, the conducting tips of which would initiate relativistic electron avalanches, whose further development is facilitated by the enhanced E field ahead of and beyond the streamer front. In other instances, TGFs appear to be initiated during brief episodes of accelerated-speed FNB.

Although obtained for downward negative breakdown of $-CG$ flashes, the results are expected to apply equally well to negative breakdown at the beginning of upward IC flashes for which the initial breakdown pulse activity is fundamentally the same as for downward CG flashes. Together, the results establish that downward TGFs of $-CG$ flashes, and satellite-detected upward TGFs of IC flashes are variants of the same phenomenon and are produced during fast negative breakdown early in the developing negative leader stage of CG and IC flashes.

2. Results

2.1. Observations

Figure 1 shows the layout of the Telescope Array Surface Detector (TASD) and the Lightning Mapping Array (LMA) used in both the earlier and present studies. The VHF interferometer (INTF) and fast electric field change antenna (FA) were located 6 km east of the TASD and utilized three receiving antennas with $106\text{--}121 \text{ m}$ baselines oriented to maximize angular resolution over the TASD (see Appendix A).

On 2 August 2018, two small, localized storms occurred over the TASD that produced three TGFs relatively close (17 km) to the INTF. The first TGF-producing discharge occurred at $14:17:20 \text{ UT}$ and was a $-CG$ flash that generated two TASD triggers $\approx 1 \text{ ms}$ after it began. The flash was initiated at $\approx 5.5 \text{ km}$ MSL altitude by a moderately high-power ($+28 \text{ dBW}$, 630 W) upward fast positive narrow bipolar event (supporting information Figure S6). The ensuing downward negative breakdown went to ground in $\approx 8 \text{ ms}$, corresponding to a stepped leader speed of $\approx 5 \times 10^5 \text{ m/s}$, somewhat faster than the normal stepped leader speeds of $1\text{--}2 \times 10^5 \text{ m/s}$. The two triggers recorded three gamma-ray bursts, jointly called TGF A, when the breakdown was at $\approx 4.5 \text{ km}$ MSL altitude (3.1 km above ground level).

Figure 2 shows “footprints” of the TA surface detections for each of the two triggered events, along with the corresponding set of scintillator observations at a central SD station. The triggers occurred within $\approx 100 \mu\text{s}$ of each other, in the southeastern corner of the TASD. The observations are similar to those reported in our previous study (Abbasi et al., 2018), in that they consisted of gamma bursts typically $10 \mu\text{s}$ or less in duration and were detected at $9\text{--}12$ adjacent SDs, over areas $\approx 3\text{--}4 \text{ km}$ in diameter. The initial burst was the most

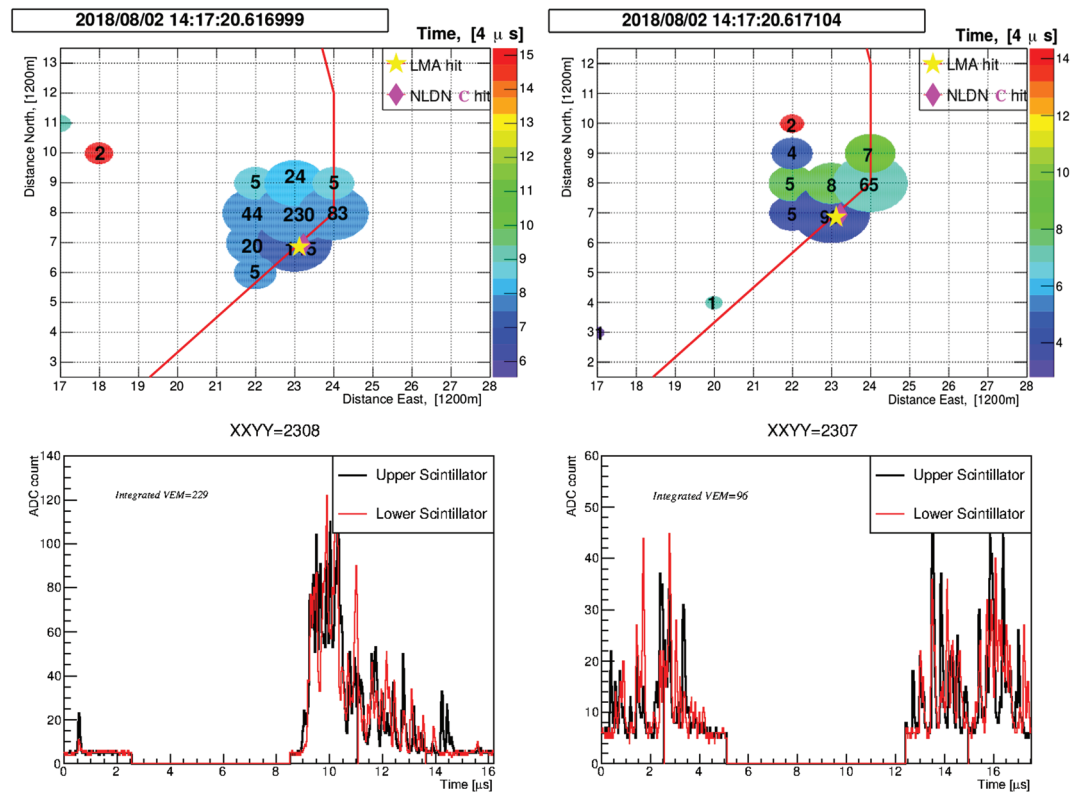


Figure 2. TASD observations of TGF A. *Top left and right:* Surface scintillator “footprints” for the three gamma-ray showers of TGF A. The grid spacing is in units of 1.2 km. The area of each circle is proportional to the logarithm of the energy deposit, and color indicates timing in $4 \mu\text{s}$ steps relative to the event trigger, corresponding to the approximate onset time of the gamma events at the ground. The yellow star shows the LMA-estimated plan location of the TGF and is in close agreement with the location of its sferic by the National Lightning Detector Network (NLDN, underlying magenta diamond) making it difficult to distinguish between the two. The red lines denote the boundary of the TASD array, showing that a portion of both showers was likely undetected. *Bottom left and right:* Scintillator responses of the surface detector stations having the largest energy deposit during each of the gamma-ray showers. The upper scintillator is represented by black traces and the lower scintillator by red traces. A single Vertical Equivalent Muon (VEM), or about 2 MeV of energy deposit, corresponds roughly to a pulse 30 ADC counts above background with 100 ns FWHM on these plots. The horizontal time axes are relative to the detectors’ individual triggers (different from the overall “event” trigger; see Appendix A1).

energetic, depositing an integrated total of 230 Vertical Equivalent Muons (VEM) (471 MeV) in the nearby TASD station, and a total of 561 VEM (1,150 MeV) over all nine adjacent stations (see Table S1).

INTF and FA observations for the flash are presented in Figure 3, which shows how the bursts were related to the discharge processes. The top panel provides an overview of the first 10 ms of the flash, from the start of the downward negative leader through the initial stroke to ground. The gamma bursts (vertical purple bars) occurred early in the flash, ≈ 1.0 and 1.1 ms after the flash’s initiation. Around this time, the FA data show a sequence of initial breakdown pulses (IBPs) of rapidly increasing and then decaying amplitude—typical of the beginning of $-CG$ flashes.

The first 1,150 MeV burst was associated with a particularly strong (-38 kA) IBP sferic, comparable in magnitude to the sferic of the ensuing return stroke, which had an NLDN-detected peak current of -37 kA. The second TGF was less strong (192 total VEM or 393 MeV) and was associated with the next-strongest IBP sferic (middle panel). Both gamma bursts were associated with episodes of accelerated downward negative breakdown.

The bottom panel of Figure 3 shows in detail how the initial gamma burst was related to the VHF radiation and sferic waveform, during a $40 \mu\text{s}$ window around the time of the burst. From the INTF elevation angles and the LMA-indicated 17 km plan distance to the source location, the VHF radiation sources descended ≈ 150 m in $10 \mu\text{s}$, corresponding to an average propagation speed $v \approx 1.5 \times 10^7$ m/s. By coincidence, this is

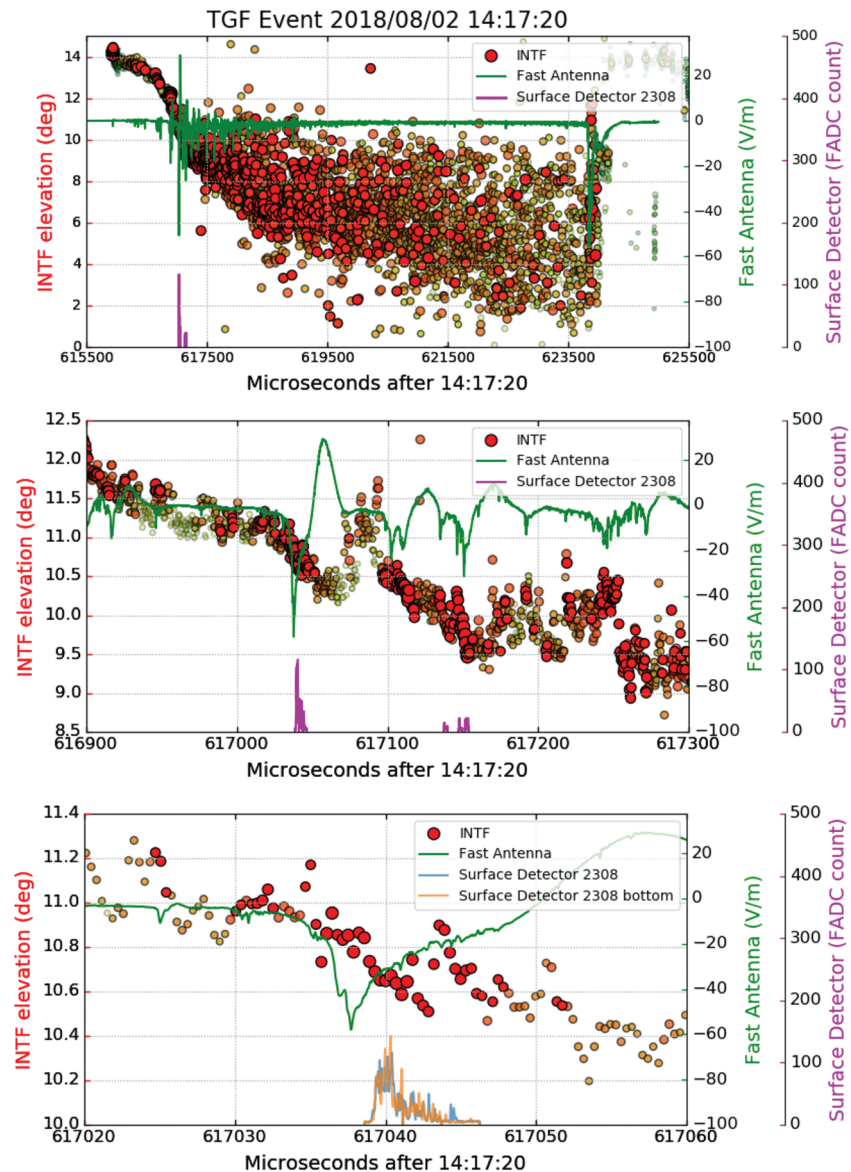


Figure 3. INTF and FA observations of TGF A. Panels show interferometer elevation versus time (circled dots, sized and colored by power), fast electric field sferic waveform (green waveform), and TASD particle surface detections (vertical purple bars). *Top:* Observations from initial breakdown through time of -38.3 kA initial cloud-to-ground stroke. Initial TGF detection occurred in coincidence with the strongest (-36.7 kA) sferic pulse, $326 \mu\text{s}$ after flash start (Table S1). *Middle:* $400 \mu\text{s}$ of observations around the time of the three gamma-ray showers of the flash, showing their correlation with the two largest amplitude initial breakdown pulses (IBPs) and episodes of fast downward negative breakdown (FNB). TASD footprints for the showers are shown in Figure 2. *Bottom:* Detailed $40 \mu\text{s}$ view of the upper and lower scintillator responses (blue and orange traces) relative to the IBP sferic and the downward FNB.

the same as the extent and speed of the upward fast positive NBE breakdown at the beginning of the flash (also ≈ 150 m in $10 \mu\text{s}$) and is indicative of the downward activity being caused by analogous fast negative breakdown (FNB) (Tilles et al., 2019). The gamma burst occurred partway through the fast downward breakdown, $\approx 1\text{--}2 \mu\text{s}$ after the peak of the negative sferic, and continued for about $5 \mu\text{s}$ before dying out shortly after the end of the FNB.

2.2. Source Determination and Time Shifting

Figure 4 shows observations of the strongest gamma-ray event for each of the TGF-producing flashes, along with time-shifted scintillator detections for each participating TASD station. The vertical line for each flash serves as a reference time for comparing the different SD waveforms with each other and with the INTF/FA.

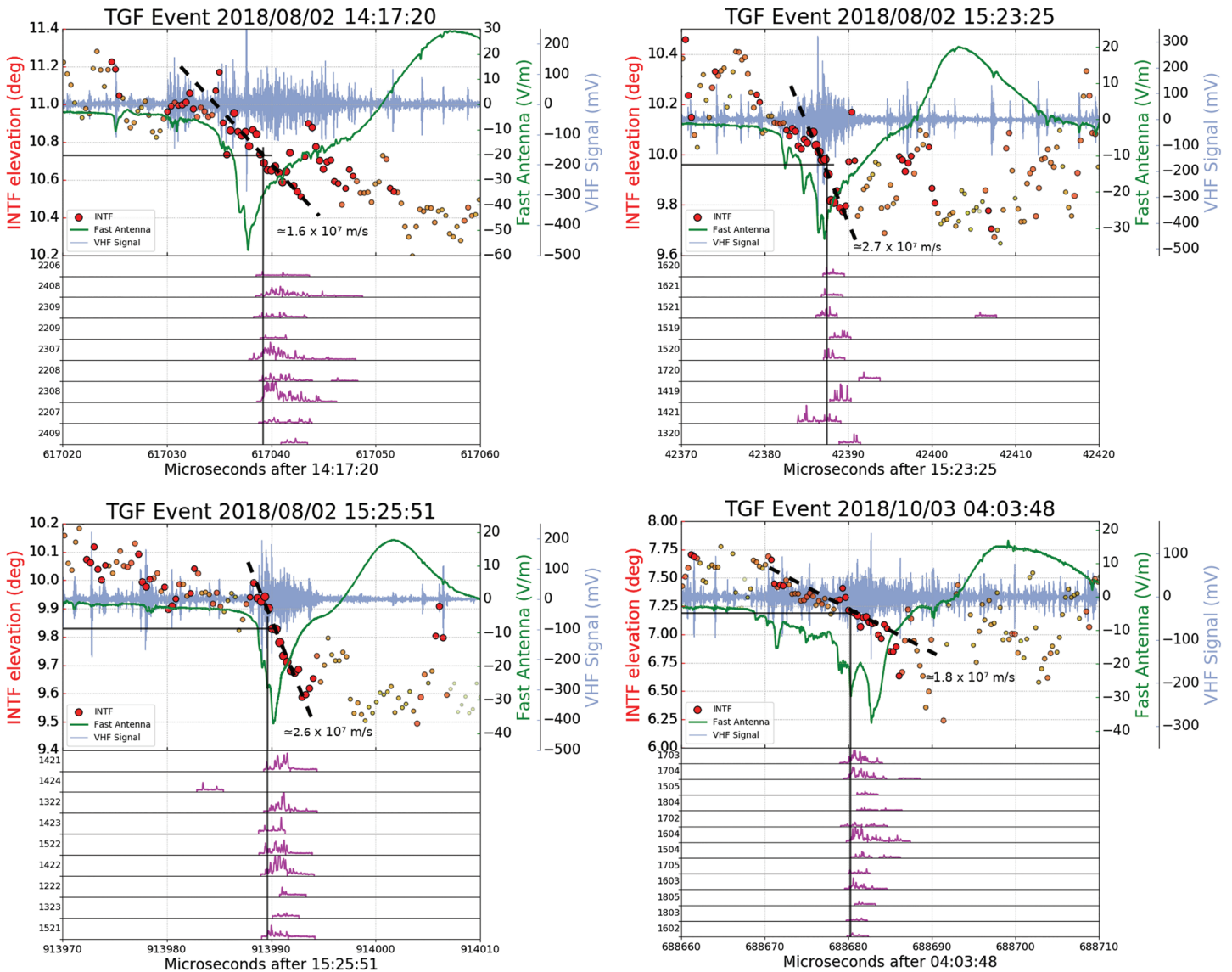


Figure 4. Detailed comparative observations. Time-shifted surface detector data for the primary gamma-ray event during each of the four TGF-producing flashes, showing how the TASD detections (lower axes) compare to each other and their relation to the VHF radiation sources and fast electric field sferics (upper axis) of the developing discharges. Black vertical and horizontal lines in each panel show the median onset time of the gamma burst(s) during the downward FNB, obtained from analysis of the collective onset times t_b at the different TASD stations and the observed INTF elevation angle versus time (see section 2.2). Light blue traces show the VHF time series waveform observed by the INTF. Station numbers XXYY in the lower axes identify each TASD's easterly (XX) and northerly (YY) location within the array in 1.2 km grid spacing units. FNB propagation speeds are indicated by the dashed lines and associated values. Full-page versions of these plots are given as Figures S15–S18.

As described below, it corresponds to the median onset time at the different SD stations. Similarly, the horizontal line indicates the elevation angle corresponding to the median source altitude immediately around that time.

The coordinate system for comparing the TASD observations with the INTF and FA data is shown in Figure A1. It is a source-centric system in which the plan position on the ground beneath the TGF serves as the coordinate origin. To shift the scintillator detection times, we need to know the slant ranges r and R from the TGF source to the SD and from the source to the INTF. The x, y plan location of the source is obtained from the LMA observations within ± 1 ms of the TGF, which determines the plan distances D and d to the INTF and to each TASD station. The TGF is therefore at point $a = [0, 0, z_a]$ in the coordinate system, where z_a is the altitude of the source above a reference plane of 1,400 m MSL. A generic TASD station is at point b , typically within ≈ 1 –3 km plan distance of the TGF. The INTF/FA is at point c , typically 15–25 km

plan distance from the TGFs. The net time shift Δt between the surface detector data at a given T ASD station and the INTF is given by the difference in propagation delays. In particular, $\Delta t = (R/c) - (r/c) = (R - r)/c$. Because the plan distances are known, the slant ranges and hence time shifts Δt are functions only of z_a . Once z_a is determined, the time shifts are calculated for each T ASD individually and used to compare the different T ASD waveforms (a) with each other and (b) with the FA spheric and the VHF source activity and centroid observations, as seen in Figure 4. For each T ASD station, the onset time at the INTF is given by $t_c = t_b + \Delta t$, where t_b is the onset time at the T ASD in question. As mentioned above, the vertical line in Figure 4 corresponds to the median of the onset times at the different stations. At the same time it also serves as a reference point for identifying stations having onset times that differ from the median value.

Because the LMA typically mislocates non-impulsive, VHF-noisy sources, the TGF's altitude is determined from the INTF elevation angles θ_c . The difficulty with doing this is that the angle changes with time during the IBP, namely, $\theta_c = \theta_c(t_c)$, making it unclear which time to pick. Even though the elevation change corresponds only to a ≈ 100 – 200 m spread in the source altitude, it corresponds to the full 10 – 20 μ s duration of the VHF and FA spheric observations. The ambiguity is resolved by recognizing that two independent measurements are necessary to determine the two unknowns, namely, the source altitude z_a and time t_a . In addition to the INTF elevation angle θ_c , the second measurement comes from onset time t_b at the particular T ASD in question. Although this provides enough information to obtain the solution, the different variables of the problem, namely, $[\theta_c, t_c, z_a, t_a]$, wind up depending upon each other, requiring an iterative approach to obtain the solution.

Figure S14 shows a block diagram of the iteration process. For each T ASD the onset time t_b is used along with an initial value of z_a to determine the corresponding onset time t_c at the INTF. The INTF data relating t_c and θ_c are then used to determine the corresponding source altitude z_a and time t_a . If the resulting z_a is different from the initially assumed value, the new value is used as the starting altitude for the next step. The iteration is stable, and convergence is reached within a couple of steps. The process is repeated for each of the participating T ASDs to obtain a set of z_a, t_a, t_c , and θ_c values from which the median is determined. Table S2 lists the full set of solutions for each T ASD of the different TGFs. The median t_c and θ_c values are shown in bold and correspond to the vertical and horizontal lines in Figure 4. For TGFs A, C, and D, the participating T ASDs all have similar onset times. The exception is TGF B, which has two or more onset times, as discussed in the next section. An analogous but somewhat different method of time-shifting and comparing the T ASD and INTF/FA observations, developed independently during the course of the study, is described in Appendix A0.2 and shown in the supporting figures. The approach utilized measurements at two T ASD stations having the strongest detections to determine the time shifts for the other T ASDs and alignment with the INTF/FA observations and provided an alternative way of investigating the observations.

2.3. Temporal Comparisons

The above analyses provide accurately determined estimates of (i) each TGF's plan location x_a, y_a , altitude z_a , and time t_a , (ii) the onset times t_c of the gamma events during the IBP, and (iii) the INTF elevation angle θ_c corresponding to t_c and z_a . The t_c and θ_c values are shown by the vertical and horizontal lines in each of the panels of Figure 4. We reemphasize the fact that the t_c values serve as reference times for comparing the different T ASD detections with each other. For TGFs A, C, and D, most or all of the stations detected the onset at the same time. The onset times are well-identified by the analysis technique and are indicative of the TGFs in question all having a single onset. An important exception is TGF B for which T ASD 1421 had a noticeably earlier onset time. Three other stations (1519, 1419, and 1320) appeared to have slightly delayed onsets. As discussed below, the different apparent onsets are notable because the footprint of the stations involved were systematically displaced in a fully 360° circular pattern around a central hole. The observations are also illustrative of the comparisons being able to identify multiple onset times.

For each of the four flashes, the gamma bursts were associated with well-defined episodes of downward-propagating fast negative breakdown. The average propagation speeds during the episodes ranged from ≈ 1.6 to 2.7×10^7 m/s (slanted dashed lines in each panel of Figure 4). This is compared to average speeds of ≈ 1.0 – 2.5×10^6 m/s for the breakdown immediately preceding the IBPs and TGFs (Figures 3 and S7–S9). The sferics associated with the TGFs constituted the strongest initial breakdown pulses of the flashes. Whereas the onset time of the gamma burst of TGF A (Figure 4a) occurred slightly after the main peak of the IBP spheric, the bursts during other flashes occurred during or at various times prior to the peak. For TGF C, the onset was at or shortly after the beginning of the IBP and FNB, while for TGF B, the primary onset was

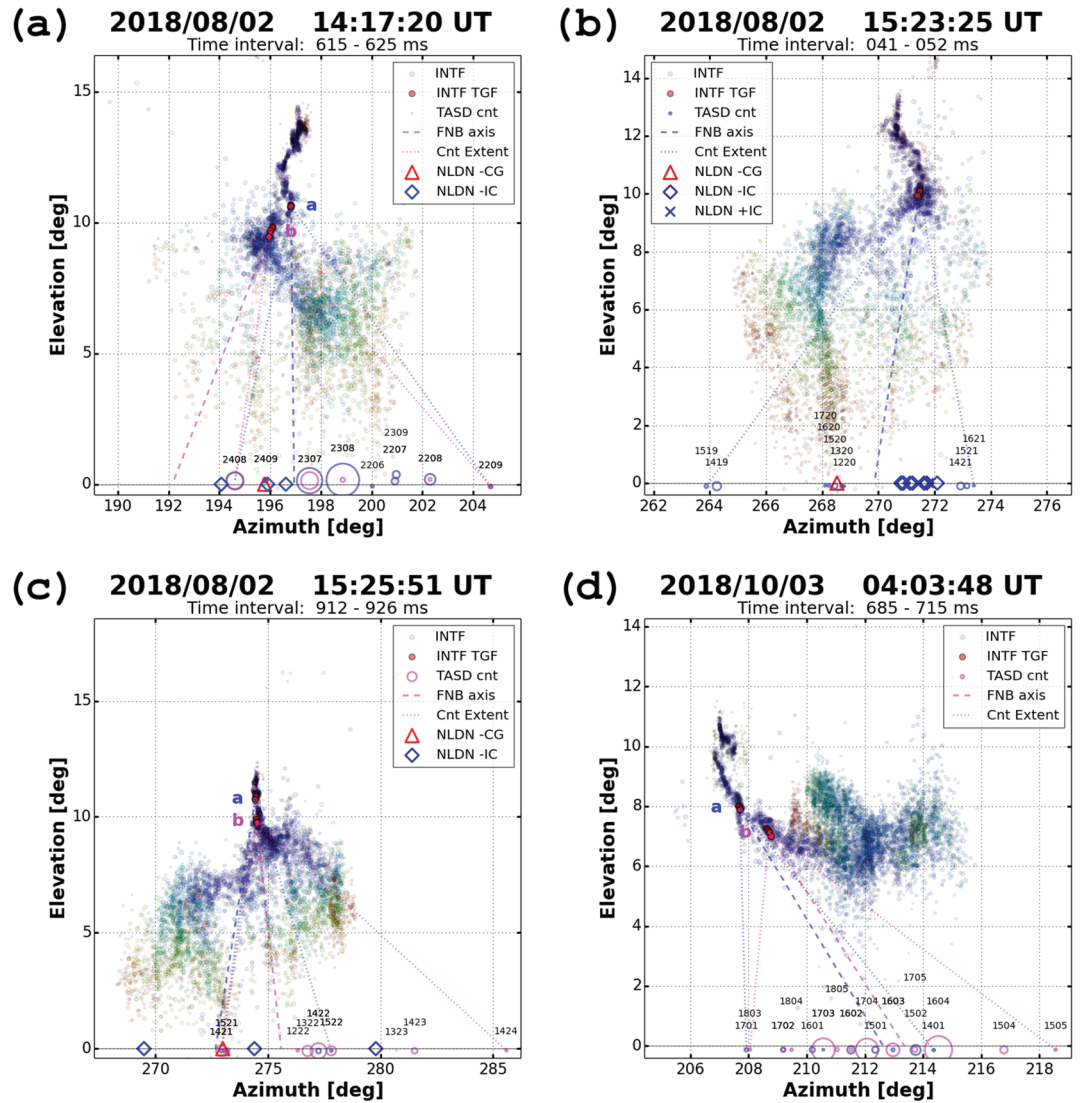


Figure 5. INTF observations of the TGF-producing flashes. Azimuth-elevation plots of INTF observations for the parent flashes of TGFs A–D, showing the initial downward development leading to the TGF occurrences (dark red sources and a and b labels, indicating the TGF altitudes). Continuation of VHF activity is shown up to the time of the initial stroke to ground for -CG flashes A–C and for a comparable time during the low-altitude intracloud flash of TGF D. Dashed lines indicate the directions of the FNB associated with each TGF and the inferred possible beaming direction. Baseline circles indicate detected TGF strength (VEM counts) and azimuthal directions of participating TASDs. Dotted line pairs indicate maximum angular spread (labelled “Cnt Extent”) of the SD detections, as viewed in the transverse plane from the INTF site. Vertical/horizontal aspect ratios are adjusted to show true angular extent. TGF B appeared to have multiple onset times at the different TASDs and therefore narrower beaming than indicated by the overall angular extent. Baseline symbols show NLDN locations of CG and IC events. Full page versions of each panel are given in Figures S19–S22.

closely correlated with the main IBP peak. For TGF D, the onset appeared to be exclusively correlated with a strong, leading-edge sub-pulse during the IBP’s FNB. IBPs having such sub-pulses are called “classic” IBPs (Karunarathne et al., 2014; Marshall et al., 2013; Nag et al., 2009; Shi et al., 2019). The sub-pulse feature of the preliminary breakdown has long been recognized, beginning with Weidman and Krider (1979), but the cause both of IBPs and their sub-pulses has remained unknown. The present results show that the IBPs are produced by fast negative breakdown and that the sub-pulses are capable of initiating gamma bursts.

For TGF A at 14:17:20 (Figures 4a and 5a), the scintillator detections in Figure 3 are from TASD 2308, corresponding to the station having the most energetic footprint. However, the estimated plan location of the burst from the LMA observations, as well as the NLDN location for the sferic associated with the burst,

indicate the breakdown was almost directly above T ASD 2307, 1.2 km to the south and 17 km southwest of the INTF (Figures S1 and S10e). The energy deposit in T ASD 2307 was slightly weaker than that in 2308 (145 vs. 230 VEM), indicating that the gamma burst was tilted slightly northward from vertical. A significant feature of the observations in Figure 4a is that the apparent onset time of the burst coincided with a step discontinuity in the VHF elevation centroid values. We later show (Figure A1b) that the discontinuity was due to a brief interval of enhanced propagation speed in which the FNB descended ≈ 50 m in $1.5 \mu\text{s}$, corresponding to a speed $v \approx 3 \times 10^7$ m/s, two times faster than the average speed of the IBP's FNB. Observations of the second set of gamma bursts during the flash show them to be similarly associated with brief episodes of enhanced fast breakdown speeds ($\approx 2.3 \times 10^7$ and 4.6×10^7 m/s; Figures S10d and S10g).

TGFs B and C (Figures 4b, 4c, 5b, and 5c) occurred in a later storm over the north-central part of the T ASD but at the same plan distances (16–17 km) from the INTF. Both were relatively weak in comparison with TGF A, with total surface detections of 112 and 212 VEM, respectively (Table S1 and Figures S2 and S3). The parent flash of TGF B was similar to that of TGF A in terms of its initiation altitude (≈ 3.9 km AGL, 5.3 km MSL) and average leader speed (1.5×10^6 m/s). The gamma bursts began 0.65 ms after flash start, again during the strongest initial breakdown pulse of the flash, whose peak current was as strong as that of TGF A (-30 kA). However, instead of the SD waveforms having a common onset time, as for TGF A, the onset times varied noticeably at different sets of T ASDs. In addition, the overall footprint of the TGF was annular-shaped around a central hole (Figure S2). The LMA and NLDN observations indicate the burst's source was over the western side of the footprint, adjacent to the hole. The initial burst was detected only at a single station, SD 1421 immediately northeast of the source. The primary onset occurred 2–3 μs later and was detected at four adjacent stations 2–3 km to the east on the opposite side of the hole (SDs 1521, 1520 and 1621, 1620). This was followed by the two southern stations having an additionally delayed onset (SDs 1519 and 1419) and finally a fourth onset back at the western-most station, almost directly below the source (SD 1720).

Concerning the correlation with the INTF and FA data for TGF B, the early gamma-ray detection at T ASD 1421 coincided with a prominent sub-pulse of the IBP and represents a separate onset time. The sub-pulse occurred during an apparently brief interlude of upward rather than downward development of the VHF radiation sources. Subsequently, the gamma-ray activity occurred during downward fast negative breakdown having a propagation speed of 2.7×10^7 m/s, with the primary onset time coinciding with the main spheric peak. Less than a microsecond after the peak, the elevation centroids exhibited a 20–30 m step discontinuity similar to that seen during TGF A, which appeared to initiate the bursts detected at the southern T ASDs.

The parent flash of TGF C occurred 2.5 min later in essentially the same location as TGF B and produced two gamma bursts 117 μs apart in time, similar to TGF A. In contrast with TGF B, both bursts were relatively simple and provide canonical examples of the basic processes of TGF production. For each event the gamma radiation was downward-directed and detected immediately below and north of the source (Figures S3 and S12). The first event was weaker and produced a total of 35 VEM (72 MeV) at four adjacent T ASDs below the source. Figure 4c focuses on the second event, which was stronger and produced a total of 212 VEM (434 MeV) at nine adjacent stations below the source. As seen in Figures 4c and S12d, the parent IBP was temporally isolated from preceding and subsequent activity, and a sudden increase of the VHF radiation signaled the onset of downward negative breakdown and the IBP spheric. The breakdown descended ≈ 120 m in $4.7 \mu\text{s}$ at a steady rate 2.6×10^7 m/s, indicative of FNB. In this simple case, the gamma radiation began immediately after the start of the FNB and continued with varying but generally increasing intensity through the entire descent until the breakdown ceased. In the process, several unresolved sub-pulses occurred, similar to the sub-pulses of TGF A. Also seen in other IBPs but more clearly shown in this flash, onset of the FNB was immediately preceded by brief upward-developing VHF sources, indicative of characteristic FPB breakdown that appeared to trigger the downward FNB.

TGF D (Figures 4d and 5d) occurred during a nocturnal storm on 3 October in a similar southward direction as TGF A but further to the south at 24 km plan distance over the southeastern corner of the T ASD (Figures S4 and S13). Again, the flash produced two triggers, the first of which contained three weak gamma bursts that were partially outside the southern boundary. The second trigger and burst occurred 140 μs later, $\approx 800 \mu\text{s}$ after the flash start. Its footprint was shifted about 2 km northward from that of the first burst, placing it entirely inside the T ASD. The apparent source of the bursts was on the eastern part of the overlapping

region between the two footprints (Figure S13e). The first burst was therefore beamed southwestward from its source, and the second burst was beamed northwestward. The westward component of the beaming is clearly evident in the INTF observations of Figure 5, which showed an increasingly strong WNW-ward tilt of the azimuthal locations as the breakdown descended, with the tilt angle becoming as large as 45° from vertical by the time of the gamma burst. A total of 440 VEM (962 MeV) was detected at 12 stations during the second burst, compared to a partial total of 100 VEM (205 MeV) at nine stations during the first burst.

Concerning the second trigger and main burst of TGF D, the IBP of the burst had a complex, relatively long-duration ($15 \mu\text{s}$) sferic waveform that was accompanied by steady downward development of the VHF radiation sources. Overall, the breakdown descended $\approx 240 \text{ m}$ in $13.4 \mu\text{s}$ at an average rate of $1.8 \times 10^7 \text{ m/s}$. The gamma burst was initiated partway through the descent, coincident with a major sub-pulse and the onset of increased VHF radiation. The sequence of events is similar to that of TGF C in that the radiation increase and corresponding sub-pulse were preceded by a brief interval of fast upward positive breakdown. The ensuing fast downward activity exhibited a small step discontinuity in the VHF centroids that coincided with the onset of the gamma burst and sub-pulse. As in each of the other TGF flashes, the gamma radiation continued up until the approximate end of the FNB, shortly after the main negative peak of the IBP sferic.

3. Discussion

3.1. Observational Results

The results of this study demonstrate that TGFs are produced during strong initial breakdown pulses (IBPs) in the beginning stages of negative-polarity breakdown. This is shown with a high degree of temporal and spatial resolution provided by a unique combination of a state-of-the-art cosmic-ray facility, coupled with high-quality VHF and LF sferic observations of the parent lightning discharges. In addition to showing how TGFs are related to IBPs, the observations reveal how the initial breakdown pulses themselves are produced, which has remained unknown for over 50 years. In particular, IBPs are produced by a recently identified type of discharge process called fast negative breakdown (FNB) (Tilles et al., 2019). FNB is the negative-polarity analog of fast positive breakdown that has been identified as the cause of high-power narrow bipolar events (NBEs) and which is instrumental in initiating lightning (Rison et al., 2016). Both polarities of fast breakdown propagate at speeds around $1/10$ the speed of light, with FPB sometimes reaching $(1/3) c$. FPB is understood to be produced by a system of propagating positive streamers that, when occurring at the beginning of a flash, is initiated by corona from ice hydrometeors in a locally strong electric field region inside storms (Attanasio et al., 2019; Rison et al., 2016).

Although the nature of fast negative breakdown is uncertain (Tilles et al., 2019), its similarities with FPB strongly suggest that FNB is also streamer-based, except for being of negative polarity. Independent of polarity or direction, both positive and negative fast streamer systems would significantly enhance the ambient electric field ahead of their advancing front (Attanasio et al., 2019), facilitating the development of high energy electron avalanches necessary for gamma-ray production.

Owing to its simplicity, TGF C provides a canonical example of the basic processes involved during an IBP. In particular, the IBP of TGF C was initiated by a brief ($1\text{--}2 \mu\text{s}$) interval of fast upward positive breakdown, immediately followed by a sudden increase in the VHF radiation and the onset of oppositely directed downward FNB (Figures 4c and S8). The positive breakdown began slightly beyond the lowest extent of the preceding negative breakdown and propagated weakly but rapidly back into preceding activity, whereupon it initiated oppositely directed and VHF-strong FNB back down and beyond the path of the upward FPB, extending the negative breakdown to lower altitude (see also Figures S12d and S12g). Similar sequences of upward positive/downward negative breakdown were associated with TGF-producing IBPs of the other flashes, including a preceding, weaker gamma-ray event of TGF C (Figures S12c and S12f).

The TGF observations show that the onset of the electron avalanching and gamma-ray production occurred at various stages during the IBPs. For TGF A, the onset occurred after the sferic peak but during still-continuing FNB. TGF C occurred at or shortly after the beginning of its IBP and FNB onset. For the more complex discharges of TGFs B and D, the onset was often associated with leading-edge sub-pulses that are a characteristic feature of classic IBPs (Karunarathne et al., 2014; Nag et al., 2009; Weidman & Krider, 1979). Like IBPs, the nature and cause of sub-pulses has continued to be a mystery (e.g., da Silva & Pasko, 2015; Stolzenburg et al., 2016). The results of the present study show that the main driving force of the IBPs is fast negative breakdown, which has the sub-pulses as embedded components. Basically, the sub-pulses are

indicative of repeated breakdown events within the developing IBP discharge. The observation that TGFs are often associated with sub-pulses, and that this occurs during fast negative streamer breakdown, provides a possible explanation for the sub-pulses' occurrence. Namely, that they are produced by spark-like transient conducting events (TCEs) embedded within the negative streamer system. That the events are spark-like is indicated by the pointed, cusp-like nature of their sferics, evidence of a sudden current onset and rapid turnoff, and also by the sub-pulses repeating several times as the IBP progresses. It should be noted that the final peak of the overall IBP sferic is also cusp-like, indicating that it too is produced by a spark-like sub-pulse.

Once initiated, the gamma radiation typically lasts $\approx 3\text{--}5\ \mu\text{s}$ for the flashes of this study. GEANT4 simulations presented in Figure S24 show that multipath Compton scattering does not artificially extend the duration, as 95% of detectable particles produced by 10 MeV (100 MeV) photons at 3 km AGL will arrive within 20 ns (60 ns). The total energy available for deposit after the first 100 ns is small enough to be indistinguishable from background levels; thus, the observed durations reflect the intrinsic duration of the sources. An important implication of this result is that relativistic avalanching lasting 3–5 μs would propagate a distance of $\approx 1\text{--}1.5\ \text{km}$, substantially beyond the 100–200 m extent of the FNB and IBP. This would provide the electron avalanches with additional amounts of electric potential energy until the ambient electric field drops below the threshold for avalanche propagation ($\approx 2 \times 10^5\ \text{V/m}$) (Dwyer, 2003).

Before proceeding, we emphasize the fact that the T ASD is detecting multi-MeV gamma radiation from the lightning discharges and not lower energy *x*-radiation. We repeat here the simple arguments for this, presented by Abbasi et al. (2018) and based on the well-understood physics of Compton electron production and the well-calibrated T ASD response to minimum-ionizing charged particles. In particular, T ASD responses for the events of the present and earlier studies (e.g., Figure S3) can clearly be resolved into individual minimum-ionizing Compton electrons that result in the deposit of approximately 2.4 MeV into either the upper or lower scintillator plane or in correlated deposits into both planes. A property of particles above the minimum-ionization threshold is that higher-energy particles would still deposit only 2.4 MeV per plane (Zyla, 2020). Thus, the T ASD cannot determine the maximum energy of Compton electrons, but it can place a lower limit on the energy values. Compton electrons that deposit 2.4 MeV into one plane are produced by a photon with no less than 2.6 MeV (Figure S9 of Abbasi 2018). Electrons that deposit 2.4 MeV into both planes, and also traverse the 1 mm steel separating sheet, have a total energy loss of 6.2 MeV and must be produced by photons with a minimum energy of 6.4 MeV.

The above inferred photon energies should be interpreted as minimal values, as they assume that the Compton electrons are produced by head-on collisions in which the gamma ray is backscattered and transfers the maximum amount of energy to the electron. The likely contributions of grazing incidence collisions to our signal would imply the actual photon energies are several times higher, depending on the grazing angle (Figure S10 of Abbasi et al., 2018). Even for single-scintillator layer detections, these are comparable to the average 7–8 MeV energy of relativistic runaway spectra detected by satellites. In any case, there is no question that the T ASDs are detecting multi-MeV gamma-rays.

3.2. Extension to Intracloud Flashes

Although obtained for downward negative breakdown at the beginning of –CG and low-altitude IC flashes, the results apply equally well to upward negative breakdown at the beginning of normal-polarity IC flashes at higher altitudes in storms. Figure 6 compares INTF and FA observations of the –CG flash of TGF C with those of an IC flash that was the next lightning discharge in the storm (see Figures S27–S29 for additional observations of the flashes). The top two panels show 2 ms of data for the two flashes with time scales of 500 μs /division. The bottom panel shows an expanded view of the large-amplitude classic IBP near the end of the IC interval. Taken together, the plots illustrate the differences and similarities of the initial breakdown processes of IC and –CG flashes. In particular, and as has long been known (e.g., Kitagawa & Brook, 1960; Weidman & Krider, 1979), the downward negative breakdown of –CG flashes intensifies more rapidly and continuously than the negative breakdown of upward IC flashes. The difference is clearly seen in the top two panels and is due to a combination of effects: first, the IC flashes needing to propagate through a relatively large vertical extent of quasi-neutral charge before reaching upper positive storm charge, compared with little or no spacing of the lower positive charge during –CG flashes (e.g., Figure 1 of Krehbiel et al., 2008, and Figure 3 of da Silva & Pasko, 2015) and secondly the IC discharges occurring at reduced pressure. The overall

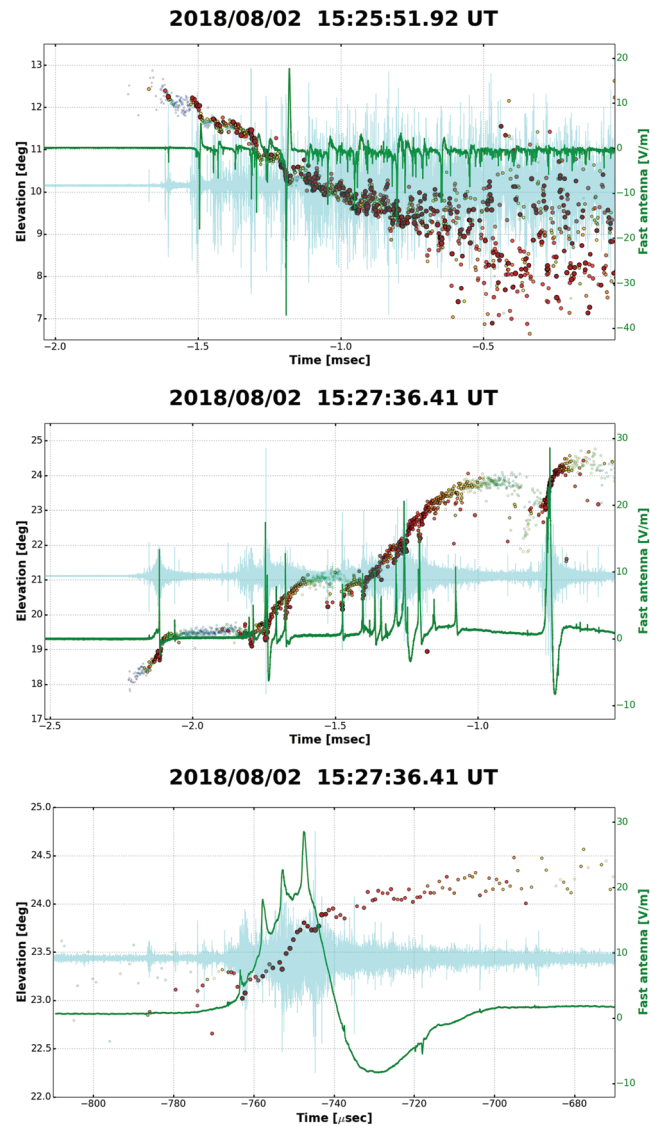


Figure 6. Comparison of the $-CG$ flash that produced TGF C with the IC flash that was the next flash in the storm, illustrating the differences and similarities between the two types of flashes. Top two panels show 2 ms of observations for the downward $-CG$ and upward IC. Bottom panel shows an expanded view of the large IBP near the end of the IC interval which, except for polarity and overall duration, is basically identical to the IBP that produced TGF B three flashes earlier in the storm. The propagation speed of the upward FNB is also similar, being $\approx 1.5 \times 10^7$ m/s.

result is that IC flashes develop more intermittently and with longer stepping lengths than $-CG$ flashes (e.g., Edens, 2014).

Despite the intensification differences, individual initial breakdown pulses of IC flashes exhibit the same features as those of $-CG$ flashes. In both instances, classic IBP sferics consist of an initial strong electric field change having embedded sub-pulses, followed by a characteristically large and relatively slow opposite-polarity field change. The similarity is illustrated by comparing an expanded plot (bottom panel of Figure 6) of the large-amplitude IBP at the end of the middle panel with that of TGF B seen in Figures 4b and S16, which occurred in the same storm ≈ 4 min earlier, three flashes before the IC flash. Except for polarity, the sferics are virtually identical. More importantly, the INTF data show both are produced in the same manner, namely, by fast negative breakdown. Owing to the increased stepping distance, IC IBPs tend to have longer durations than those of $-CG$ s, lasting $\approx 70 \mu\text{s}$ for the IC IBP versus $\approx 35 \mu\text{s}$ for the IBP of TGF B. The fast negative breakdown component of the IC IBPs is also similarly longer, being $\approx 20 \mu\text{s}$ for the IC versus $\approx 10 \mu\text{s}$ for TGF B. The factor of two overall duration difference agrees with the study by Smith

et al. (2018) of median durations of large IBP spherics in Florida storms. Another example of a similar classic IC IBP spheric is seen in Figure 4 of the study of Florida IBPs by Marshall et al. (2013), which had a duration of $\approx 100 \mu\text{s}$ and was considered to be a “candidate” TGF flash. At this point it should be noted that in many instances the durations of IC and CG IBPs are the same for both types of flashes. This is seen in the scatter diagram of Figure 5 of Smith et al. (2018) and is shown in detail by the comprehensive observations of Tilles (2020). Figures 9.3 and 9.4 of the latter study, conducted in Florida with the same INTF and FA instrumentation as in the present Utah study, show that (except for polarity) the IC and -CG IBPs were essentially indistinguishable both in terms of their spherics and durations.

Due to the TGF-producing storms having low flashing rates (typically 1–2 min between flashes in the present study), the electrification is allowed to build up to large values, causing both the -CG and IC flashes to be highly energetic when they finally occur. For the IC flash of Figure 6, this is reflected not only in the amplitude and duration of the classic IBP but also by the preceding activity being produced by two complex sequences (clusters) of IBPs and sub-pulses, seen in the middle panel. Each of the clusters is linked together by continuous, upward-developing high power negative breakdown, producing long-duration complex steps. The overall durations of the two clusters were ≈ 130 and $400 \mu\text{s}$, respectively. Expanded views of the complex IBPs are seen in Figure 7, which show the spherics were dominated by increasing numbers of sub-pulses that assisted in continuing the negative breakdown and extending the cluster durations. In addition to their increased numbers, the sub-pulses are dramatically more impulsive and stronger in amplitude than those of the -CG flashes. The IC sub-pulses had amplitudes of ≈ 10 – 20 V/m , compared to ≈ 5 – 10 V/m for the sub-pulses (at essentially the same distances) of the TGF-producing IBPs of Figure 4 (seen in larger scale in Figures S15–S18). Given that the simpler IBPs of the -CG flashes produced TGFs, the IC flash would likely have been equally or more capable of generating upward TGFs. Due to relativistic avalanching being a strong function of the potential difference being shorted out by the spark-like sub-pulses (Celestin et al., 2015), as well as the sub-pulses being more dynamic (Celestin et al., 2012) and repetitively impulsive, the resulting avalanching and TGFs would be more energetic, as well as longer lasting. Similar observations were obtained for an IC flash that occurred between TGFs B and C, which are compared with TGF B in Figures S24–S26.

3.3. Implications for TGF Production Mechanisms

As summarized in the recent modeling study of TGFs by Mailyan et al. (2019), there are two classes of models for TGF production: First, what is termed the relativistic runaway electron avalanche (RREA) or relativistic feedback (RFD) model in which electron avalanches develop in km-scale regions of strong electric fields in storms (Dwyer, 2003). In this model, the avalanching is enhanced by relativistic feedback that increases the avalanche currents by several orders of magnitude (Dwyer, 2012). The second class is broadly termed the “leader” model in which the relativistic avalanches are initiated in the highly concentrated electric field produced at the negative tip of a conducting leader channel. The electric field at the tip is extremely strong as a result of the leader having kilometer-scale extents and shorting out tens to a few hundred MV of potential difference in the storm. Whereas the RREA process by itself requires cosmic ray-produced or other seed relativistic electrons to get started, the leader process begins with low energy thermal electrons and requires exceedingly large electric fields ($\approx 3 \times 10^7 \text{ V/m}$ —an order of magnitude larger than the breakdown strength of air) to be accelerated into the runaway electron regime, where their number and energy increase exponentially with time and distance (e.g., Dwyer, 2004). Electric fields of this strength are produced only at the tips of conducting leader-type channels and then only transiently during rapid channel development. Thermal electrons are accelerated into the relativistic regime as a result of transient negative streamers within the strong E region (the so-called “negative corona flash”), as described by Moss et al. (2006), Celestin and Pasko (2011), and Celestin et al. (2015). Once the leader/streamer-initiated avalanches are started they would be able to initiate the relativistic feedback process.

While relativistic feedback can explain the large currents and fluxes of highly energetic satellite-detected events, it does not appear to be playing a role in initiating the smaller-scale observations of the present study. Instead, the inference that IBP sub-pulses is caused by spark-like transient discharges embedded within the fast negative streamer system points to the leader/streamer model as playing an important and possibly dominant role in generating runaway avalanches and TGFs. Once initiated, the runaway electrons would additionally increase in energy while propagating through the enhanced field region ahead of and beyond the relatively broad streamer front (Attanasio et al., 2019).

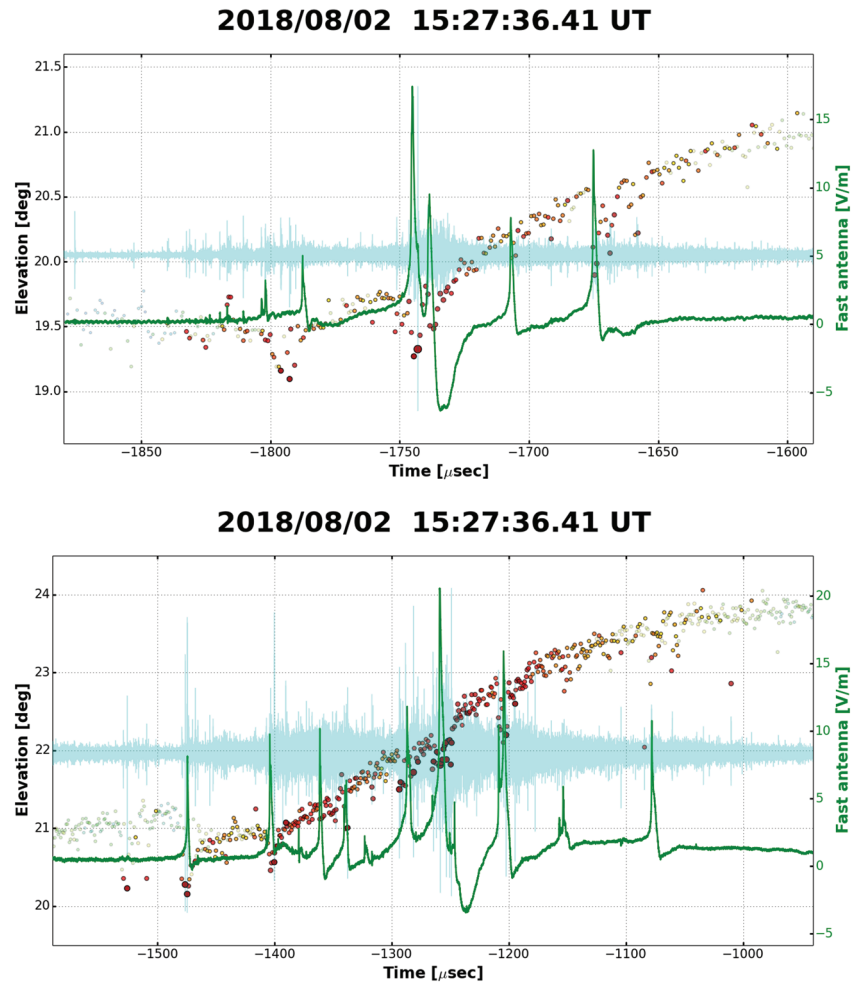


Figure 7. Expanded views of the complex IBP clusters of the IC flash of Figure 6, showing the increased number and highly impulsive nature of the sub-pulses. The FNB breakdown of the IBPs and the sub-pulses are each embedded in continuous upward negative streamer breakdown having a propagation speed of $\approx 2\text{--}3 \times 10^6$ m/s, showing that negative streamer breakdown doesn't have to travel at speeds of 10^7 m/s to produce the sub-pulse sparks. The durations of the two clusters were ≈ 130 and $400 \mu\text{s}$, respectively, with the spheric of the first cluster resembling that of the TGF-producing IBP of Figure 2 of Lyu et al. (2018) and Pu et al. (2019) and the second cluster resembling the spheric of another complex TGF-producing spheric of Pu et al.

An important question is whether the conducting channels of the sub-pulses (which we refer to as transient conducting events or TCEs) are isolated within the negative streamer system and from each other or if they are connected back into, or originate from, the conducting channel of the incoming negative leader. If so connected, the potential drop beyond the negative tip of the sub-pulse channel would be comparable to the amount shorted out by the km-long or longer leader, envisioned to be as large as 60 to 200 MV or more (e.g., Celestin et al., 2015; Mailyan et al., 2019). Such a leader is termed a “high potential” leader, which by itself can produce the large ($\approx 10^{16}\text{--}10^{18}$) gamma photon fluxes inferred by satellite observations (Celestin et al., 2015).

To address the question of the sub-pulse connectivity, we note that the sub-pulses continue to occur until one suddenly causes the IBP spheric to begin transitioning to an opposite-polarity field change during the final part of the IBP. Although the flash current does not change direction, the electric field waveform becomes dominated by the electrostatic and induction components, which are inverted in polarity from the radiation component due to the flash being beyond the reversal distance d for vertical dipolar discharges, where $d = \sqrt{2}h$ and h is the discharge height above ground level (e.g., MacGorman & Rust, 1998). At the same time, the fast negative breakdown continues to propagate for several microseconds before finally dying out. From the large amplitude and relatively long duration of the opposite-polarity field change, one can infer

that the current is not constrained to the IBP itself but develops retrogressively back through the negative breakdown leading to the IBP, converting a potentially weak streamer-leader channel to a hot conducting leader and completing the step. That the current during a negative leader step develops in a retrograde manner back along the incoming breakdown channel has been shown by in situ balloon-borne observations of negative leader stepping during an IC flash by Winn et al. (2011) and by high speed video observations around the time of IBPs of -CG flashes by Stolzenburg et al. (2013), as discussed later. Because sub-pulses previous to the final sub-pulse do not initiate the opposite polarity field change, one can infer they are not connected to the incoming leader breakdown but instead are isolated from the leader and from each other. The question then becomes whether the sub-pulse sparks short out enough potential difference to account for the observed TGFs.

In terms of the space stem/space leader model of negative leader stepping (e.g., Biagi et al., 2010; Petersen et al., 2008), the sub-pulse sparks would correspond to conducting space leaders that occur in the negative streamer region ahead of the developing leader. Continuing the space leader interpretation, the final sub-pulse develops back into the incoming leader at which point the leader's potential rapidly advances to the opposite end of the space leader, producing the negative corona flash that launches the relativistic electrons. This scenario could explain TGF A, which was initiated a few microseconds after the final, sharply pointed negative peak of the sferic (Figures 4a and S15). TGF A also produced the most surface-detected energy of the different TGFs (561 VEM total or 1150 MeV; Table S1). Because the TGF occurred just above the T ASD boundary (Figures 2 and S1), the detected energy could have been up to 50% larger had it been entirely captured. Similarly, the scenario could also explain the main onset of TGF B, which occurred at the same time as the final sub-pulse peak (solid vertical line in Figures 4b and S16).

For TGFs C and D, however, and for the early initial detection of TGF B, the TGF onsets were associated with sub-pulses that did not initiate a retrograde current (Figures 4, S17, S18, and the left-most vertical dotted line in Figure S16). These and the other early sub-pulses of the IBPs would be characterized as attempted space leaders and may have somehow paved the way for the final sub-pulse but otherwise appeared to be independent of each other and not connected back to an incoming leader. The gamma events of TGFs C and D had total surface detections of 212 and 440 VEM (434 and 902 MeV), respectively, with TGF D being the second strongest TGF after TGF A. At the same time, the total activity of TGF B, which was most closely associated with the IBP's final sub-pulse and presumably the best candidate for being connected to the incoming leader, had the weakest total surface detection of all, 112 VEM (229 MeV).

Storm-to-storm variability, as well as that from flash to flash in the same storm, coupled with the small sample size makes it difficult to compare the different observations. However, the fact that three TGF events (C, D, and the initial lone detection of TGF B) were initiated by sub-pulses that did not connect back into the incoming breakdown of the IBP, and the subsequent activity of TGF B producing a weak TGF despite its sub-pulse eventually connecting back into the incoming breakdown, indicates that the occurrence and strength of the gamma bursts are determined more by the amplitude and impulsiveness of the initiating sub-pulse rather than by the incoming breakdown consisting of a hot conducting leader.

From the above results, as well as the IBPs being produced by fast negative streamer breakdown, the sub-pulses are analogous to the space leader in negative leader stepping in that they occur within negative streamers ahead of the leader. Instead of being produced by a relatively slow-developing thermal space stem, the sub-pulses are impulsive sparks caused by sudden instabilities in extended-length streamer channels associated with fast propagation speed of streamers. And instead of the impulsivity of the step being produced by the space leader suddenly contacting a conducting leader channel and rapidly propagating the leader potential forward to the head of the space leader, the impulsiveness and negative corona burst are produced by the spark itself. The succession of sub-pulse sparks eventually causes one to develop back into a somewhat diffuse leader, giving rise to the backward-developing current that further establishes and converts the incoming breakdown into a well-defined hot conducting channel. This scenario agrees with high-speed video observations by Stolzenburg et al. (2013, 2014), indicating that the "unusual" steps of IBPs occur ahead of a weakly-conducting nascent leader rather than a continuously hot, conducting channel (see later discussion).

If the space stem/space leader process is what initially advances the conducting leader channel, a legitimate question concerns how such a hot leader is produced in propagating from the end of the preceding IBP

(or from the flash start) to the beginning of the IBP in question, in the absence of discernible space stem/space leader activity. At some point the leader becomes self-propagating (e.g., da Silva et al., 2019), but apparently, this does not occur in the early stages of the breakdown, as evidenced by the increasing need for and strength of IBPs in the initial few milliseconds of negative breakdown. Up until then, the advancing negative breakdown between IBPs appears to be a system of relatively weakly conducting negative streamers, which can self-propagate more readily.

From the INTF observations, the average speed of the downward negative breakdown at the beginning of the TGF-producing flashes is $\approx 1.0\text{--}2.5 \times 10^6$ m/s (e.g., Figures 3a and S7–S13), an order of magnitude or so faster than other estimates of developing leader speeds (e.g., Behnke et al., 2005). Similarly fast progression speeds were reported during the upward development of TGF-producing IC discharges by Cummer et al. (2015), who used ionospheric reflections to determine the altitude and hence the upward progression speed of successive radio pulses of TGF-producing IC flashes. For three different flashes, the speeds were noted to be remarkably similar and fast, ranging from 0.8 to 1.0×10^6 m/s. As in the present study, the TGFs were produced partway along the vertical development (in their case upward), when the leader was $\approx 1\text{--}2$ km in extent. The fact that TGFs were not also produced by subsequent pulses at higher altitude during the vertical development led them to ask why this did not happen, in view of the leader lengths being proportionally longer. A similar question would apply to the present, downward-directed observations at the beginning of the –CGs.

Taken together, the results suggest a scenario in which a “step” consists of (a) intermediate-speed negative streamer breakdown being launched at the end of the previous step’s IBP, which progresses in a forward direction until (b) initiating accelerated-speed FNB and an IBP having embedded sub-pulses, one of which (c) initiates a strong current that develops retrogressively backward through the IBP and its preceding negative breakdown, thermalizing and extending the negative leader. The IBP then reverts back to intermediate or slower-speed negative streamer breakdown, beginning the next step. Whether a TGF is produced during the IBP is largely decoupled from the preceding negative breakdown, explaining the independence of TGF production on the extent of the negative breakdown up to that point. Where the preceding extent plays a role is in enhancing the electric field ahead of its developing front, to the point that the FNB is initiated. The field enhancement is due to the cumulative dipolar charge transfer of the negative breakdown during each step (e.g., Attanasio et al., 2019; Cummer, 2020; Krehbiel, 2018), causing successive IBPs to become stronger with time. The TGFs of this study were produced by the strongest IBP of the flash, but in three of the four flashes one or two additional bursts occurred that were associated with separate episodes of FNB and sub-pulse activity (see Figures S10d, S10g, S12c, S12f, S13c, and S13f). The additional gamma events occurred during less strong IBPs within $\approx 100\text{--}150$ μs either before or after the main gamma events and represent sparsified examples of the TGF activity that would be expected during the kind of complex IC IBP events seen in Figures 6 and 7.

The above scenario for the stepping provides an explanation for the optical observations of Stolzenburg et al. (2013) in which partially obscured luminosity in the first 1–2 ms of a –CG flash advanced downward with a series of surges associated with bright optical emissions at the times of successive IBPs. The observations were obtained from high speed video recordings having 20 μs time resolution. Each bright surge lasted about 80–100 μs and was preceded by dim, linearly downward extension of the channel, with the brightest frame “immediately followed by backward lighting of the entire tail” that preceded the bright surge. The sequence then started over again with renewed dim downward extension of the channel to a lower elevation angle, with the process repeating for up to five surges. In terms of the above scenario, (a) the linear downward channel extensions would correspond to the intermediate-speed, inter-IBP negative streamer activity, (b) the succeeding bright optical emissions would have been produced by the spark-like sub-pulses of the IBP, and (c) the immediately following upward propagating light would be produced by the retrograde current traveling back up along the path of the pre-IBP activity, converting it into a hot conducting leader. As noted earlier, Winn et al. (2011) observed similar backward propagating current events following individual steps of an already developed negative leader toward the end an IC flash, using close balloon-borne electric field change observations of the flash. The correlation of bright optical pulses with –CG IBPs was extended

by Stolzenburg et al. (2016) to be produced by IC-type IBPs at the beginning of hybrid –CG flashes. Similar to Marshall et al. (2013), the IBPs were considered to be candidate producers of TGFs, on the basis of the IBPs being complex and having strong sub-pulses.

The mechanism for producing the spark-like sub-pulses and TCEs within the fast negative breakdown would be essentially the same as that which causes the FPB and FNB to be the producer of high-power VHF radiation, described as being the strongest natural source of VHF radiation on Earth (LeVine, 1980). Due to their fast propagation speed, both polarities of streamers would have extended partially conducting tails that would become unstable in the strong ambient fields (Malagon-Romero & Luque, 2019; Shi et al., 2016). The resulting rapid current cutoff, coupled with meters-long extents and large numbers, make both polarities of streamer systems potent radiators at VHF (Rison et al., 2016). The negative polarity streamers of FNB would have more robust and extensive tails than positive streamers that could occasionally extend over longer distances, with the resulting instabilities and currents producing hot, spark-like conducting channels of the sub-pulse TCEs. In addition to explaining the optical emissions associated with IBPs, the sudden occurrence of a dynamically impulsive conducting channel would provide the means for initiating relativistic electron avalanches (Celestin & Pasko, 2012; Celestin et al., 2015; Moss et al., 2006).

It is interesting to note that, in addition to being produced by sub-pulses, it may also be possible for relativistic electron avalanches to be initiated by individual negative streamers themselves. This is suggested by the modeling study of Moss et al. (2006), who showed that the extremely strong electric fields sufficient to accelerate electrons into the runaway regime will occur briefly immediately prior to branching of the streamers. Electrons produced in association with branching can reach kinetic energies as large as 2–8 keV or larger, well into the runaway electron regime. Although determined to occur in the corona flash and streamer zone at the tip of a conducting leader, the process might also occur at the tips of streamers having relatively long conducting tails. The branching process was noted to strongly favor negative streamers over positive, due to positive streamers requiring photoionization to sustain their propagation. If it occurs, the branching mechanism would be a powerful adjunct to TCEs, since large numbers of individual streamers exist within a propagating system that are spread over a much larger cross-sectional area than an individual conducting leader or TCE channel and are continually branching.

Other issues of note concerning the observations are (a) that the TGFs are broadly rather than narrowly beamed, favoring a tip-based conducting channel model (Mailyan et al., 2019) and (b) are commonly tilted at substantial angles from vertical. From the TASF footprints and source altitudes, the half angular width of the beaming is on the order of 35° or so (≈ 2.4 km radial plan spread for a 3.3 km source altitude). From the INTF observations of Figure 5 (repeated in larger scale in Figures S19–S23), the tilting can be 45° or more, depending on the three-dimensional development of the discharge. Finally, successive sub-pulses can be oriented in different directions, as indicated by successive onsets occurring in different directions for TGF B (Figures S16 and S20).

We note that the simulations of our previous study (Abbasi et al., 2018) implied TGF fluences on the order of 10^{12} – 10^{14} relativistically generated gamma photons, several orders of magnitude less than satellite-inferred fluences of $\approx 10^{16}$ – 10^{18} photons. From Celestin et al. (2015) (Table 1), total fluences of 10^{12} – 10^{14} photons correspond to potential drops of ≈ 10 –50 MV or so at the conducting channel tips, while fluences of $\approx 10^{16}$ – 10^{18} photons correspond to larger potential drops of 160–300 MV. That the observed fluences are relatively weak would be consistent with the inference that the TGFs are produced by isolated conducting sparks that short out lesser amounts of potential difference. However, if km-long conducting leaders are not involved, the question is whether sufficient potential difference is available for producing the relativistic electrons and the observed gamma radiation. For example, from Celestin et al. (2015) (Figure 3), 5–10 MV potential drops would not produce relativistic electrons greater than ≈ 1 –2 MeV. On the other hand, 60 MV (160 MV) of potential drop would produce relativistic electrons up to 9 MeV (20 MeV). From the modeling, then, at least 50 MV of potential drop would be required to produce the expected gamma energies observed in this study. The predicted fluences corresponding to 60 MV (160 MV) potential drop, however, is $\approx 6 \times 10^{14}$ ($\approx 4 \times 10^{16}$) photons, two orders of magnitude greater than the inferred fluences of these TGFs. Thus, the observations are inconsistent with the leader-streamer modeling, in that the fluences corresponding even to the minimum likely detected photon energy produced by 60 MV potential drop would be at the upper end of the implied fluence values of Abbasi et al. (2018).

The question of available potential energy can be addressed by considering the electric field required for streamer propagation, called the stability field E_{st} . From da Silva and Pasko (2013), at one atmosphere of pressure $E_{st} \simeq 5 \times 10^5$ V/m for positive streamers, but $\simeq 12.5 \times 10^5$ V/m for negative streamers in virgin air. The fields scale according to pressure, so at 5 km altitude (0.5 atm) $E_{st}^- \simeq 6 \times 10^5$ V/m. Thus FNB propagating over the 100–240 m long extents of the TGF IBPs (Table S3) would experience total potential differences of $\simeq 60$ –150 MV, with 60 MV being consistent with observed photon energies up to $\simeq 9$ MeV. Some or all of the potential difference that is not shorted out by the sparking would be available for additional avalanche growth down to the propagation threshold of 2×10^5 V/m, which is not accounted for in the Celestin et al. (2015) calculations. Also not accounted for are dynamical effects in initiating the relativistic electrons that are associated with the sparking being impulsive, which are significant for pulsed discharges (section 5.4.3 of Nijdam et al., 2020). Finally, using the stability field values doesn't account for the field intensification ahead of the advancing streamer front, which can be as much as 50% above the ambient E_{st} value (e.g., Attanasio et al., 2019; da Silva et al., 2019). For IC flashes at higher altitudes, E_{st} would be reduced by about another factor of two, but this would be offset by the IC events typically being longer by a factor of two or more, leaving the total potential differences about the same. Finally, we note that vertical profiles of the electric potential in electrified storms similar to those being studied show the total potential differences available for IC and –CG flashes are both on the order of 200 MV (e.g., Figure 1 of Krehbiel et al., 2008; Figure 3 of da Silva & Pasko, 2015).

In short, while the details remain to be understood, taken together, sufficient potential difference is available to produce gamma radiation into the 10–20 MeV range or potentially higher, consistent with the observations and the physics of the Surface Detector responses. The main issue is the fluence values. A possible explanation for the fluence inconsistency that allows both the observational data and the modeling to be correct would be that the gamma photons are produced by $\simeq 10$ –50 MV of potential drop, which from Figure 3 and Table 1 of Celestin et al. (2015) would produce relativistic electron energies in the range of $\simeq 2$ –9 MeV and fluences in the observed range of 10^{12} – 10^{14} photons. Once initiated, the electron energies would be further accelerated up to $\simeq 10$ –20 MeV by the enhanced field ahead of the streamer front and any ambient field beyond greater than the threshold field of 2×10^5 V/m. Because the extent of the field ahead of the streamer system would be less than an e-folding avalanche length, the fluences would not change significantly while the electron energies increase.

To the extent that satellite-detected TGFs from IC flashes have substantially larger fluences, the implication is either (a) that the satellite detected events emanate from the tips of fully formed, kilometer-length or longer conducting leaders in which case fluences of 10^{16} – 10^{18} photons are achieved directly from the negative corona flash produced by potential drops as large as several hundred MV or (b) that the fluences of lesser potential drops are enhanced by the relativistic feedback process. The abovementioned observations by Cummer et al. (2015) raise the important question about the leader hypothesis of why TGFs are not produced later in the development of upward, kilometer or multi-km conducting leaders. Instead, and as additionally discussed below, the observational data support the idea that the much greater satellite-detected fluences are due to the relativistic feedback mechanism, which was initially developed to explain this very issue (Dwyer, 2012).

Another substantial difference between the present observations and those obtained by satellites concerns the durations of the TGFs, being 5–10 μ s for the downward –CG TGFs, versus $\simeq 20$ –200 μ s for the upward, IC-generated TGFs (e.g., Mailyan et al., 2016, 2018; Østgaard et al., 2019). The difference can be at least partially explained by observations that IC flashes can often have long-duration, complex sferics, consisting of multiple sub-pulses and IBPs, each of which would be capable of producing TGFs. Examples of such sferics are seen in Figures 6 and 7. Of particular note are the observations of three TGF events by Lyu et al. (2018) in which complex dB/dt events produced Fermi-detected TGFs having continuous durations of $\simeq 50$, 100, and 120 μ s. In the latter two cases, gamma detections occurred intermittently for an additional 60 and 100 μ s both before and/or after the main activity, extending their overall durations to $\simeq 160$ and 220 μ s, respectively. For each of the three events, the TGFs were produced during the occurrence of a slow, smooth component of the sferic, indicative of being caused by electron avalanching that produced the TGFs. Complex, lengthy sferics were also produced by the other two events of the same Lyu et al. study.

Of particular interest, and the best-studied example, was the first event of 4 September 2015 (Figure 2 of Lyu et al., 2018), which occurred over west-central Florida. Its sferic closely resembled that of the first complex

IBP of the Utah IC, seen in the top panel of Figure 7. In both cases, the sferic lasted for $\approx 250 \mu\text{s}$ and consisted of several highly impulsive sub-pulses before and after a central event. For the Utah IC the central event was itself a large-amplitude IBP, while for the Florida IC it was the large-amplitude slow field change of the electron avalanche. The comparison, along with the other Lyu et al. examples illustrates the fact that (a) long-duration TGFs can be produced by IC flashes having complex sferics and (b) that the only difference between the Utah and Florida ICs is that the latter initiated strong runaway avalanching, while the former did not but based on the sferic similarities could well have done so. The second complex IBP of the Utah IC, seen in the bottom panel of Figure 7, would have been even more capable of generating a long-duration TGF based on its greater duration and VHF signal strength.

Pu et al. (2019) extended Lyu et al.'s study to include five additional examples of continuously and intermittently long-duration TGFs being produced by other IC flashes having complex IBP sferics. Finally, we call attention to the study by Tilles et al. (2020) of a high peak current (247 kA) energetic in-cloud pulse (EIP) that was observed in Florida with the same physical INTF and FA instrumentation of the present study. The EIP was produced by a complex sequence of repeated IBP-type fast breakdown activity, but its sferic was completely dominated by a sequence of three successive slow, smooth relativistic avalanches indicative of being produced by relativistic feedback. No gamma-detecting satellite happened to be in view of the EIP, but the flash undoubtedly produced an upward TGF (Cummer et al., 2017; Lyu et al., 2016) and is an example of how IC flashes are capable of producing extremely strong avalanching as a result of complex IBP-type activity.

3.4. Summary

The results can be summarized as follows:

1. Downward TGFs occur during strong, "classic" initial breakdown pulses (IBPs) of downward negative CG and IC flashes. In turn, the IBPs are produced by streamer-based fast negative breakdown (FNB).
2. The TGFs consist of short, $\approx 5\text{--}10 \mu\text{s}$ duration bursts of gamma rays initiated by sub-pulses during the IBPs and apparently also by brief episodes of enhanced speed FNB.
3. The correspondence of TGFs with sub-pulses is indicative of the sub-pulses being produced by spark-like transient conducting events (TCEs), consistent with their sferics being impulsive or cusp-like and explaining the bright optical activity observed during IBPs of $-CG$ and IC flashes.
4. In turn, the TCEs are considered to result from instabilities in occasionally long streamer tails or partially conducting channels embedded within the FNB of the IBP and to be isolated from each other and from the incoming breakdown preceding the IBP.
5. Based solely on the well-understood physics of surface detector responses and Compton electron production, individual electrons detected by the T ASD surface stations correspond to photon energies no less than 2.6 MeV if detected in a single scintillator layer and 6.2 MeV if detected in both layers.
6. From the electric field required to propagate negative streamers in virgin air at $-CG$ altitudes, the electric potential difference experienced by the FNB over the 100–240 m extents of TGF-producing IBPs is $\approx 60\text{--}150$ MV.
7. Instead of the breakdown leading up to an IBP being a long conducting leader, it appears to be due to weakly conducting negative streamer breakdown that gets accelerated to produce the IBP.
8. The observational data indicate that the streamer to leader transition of successive steps is caused by current generated during the characteristic opposite-polarity field change in the final stage of the step's IBP.
9. The initial upward negative breakdown of IC flashes is shown to be produced in the same basic manner as the initial downward breakdown of $-CG$ discharges but generally lasting longer and having longer step sizes.
10. The long durations of satellite-detected TGFs can be explained by IC flashes producing complex clusters of sub-pulses and IBPs, which enable the development of continuous and intermittent electron avalanching. Sparse versions of this are seen during successive IBPs of $-CG$ flashes.

While the present study has been underway, the T ASD has been in the process of expanding by a factor of four in its coverage area, and the TGF and lightning observations are continuing. The LMA network is being similarly expanded, and an additional VHF interferometer instrument is to be added in the current year. Detailed analyses of additional observations are the subject of continued study.

Appendix A: Methods

A1. Instruments

Telescope Array Surface Detector. The TASD consists of 507 scintillator detectors arranged on a 1.2 km square grid. The array is situated on a relatively high, 1,400 m altitude desert plain in west-central Utah, and covers an area of $\approx 700 \text{ km}^2$. Each detector has two scintillator planes, each $3 \text{ m}^2 \times 1.2 \text{ cm}$ thick, separated by a 1 mm thick steel sheet and housed inside an RF-sealed and light-tight stainless steel enclosure. The TASD is designed to detect the charged components—primarily electrons, positrons, and muons—of the cosmic ray-induced Extensive Air Shower (EAS). An event trigger is recorded when three adjacent SDs observe a signal greater than that of 3 Minimum Ionizing Particles (MIPs) (≈ 150 FADC counts) within $8 \mu\text{s}$. When an event trigger occurs, the signals from all individually triggered SDs within $\pm 32 \mu\text{s}$ are recorded (Abu-Zayyad et al., 2013). An individual SD trigger occurs upon observing a signal of amplitude greater than 0.3 MIP (≈ 15 FADC counts) within $8 \mu\text{s}$.

The TASD is an inefficient detector of gamma radiation, relying on the production of high-energy electrons through the Compton scattering mechanism in either the thin scintillator, steel housing, or air above the detector units. Detailed simulations of this process have been described in the authors' previous study (Abbasi et al., 2018). Incident gamma-ray photons with energy above 10 MeV will on average deposit about 20% (30%) of the energy of a MIP in the upper (lower) scintillator. The majority of photons will not interact in the detector at all; those that do will primarily create Compton recoil electrons with kinetic energies at or below the photon energy level. The Compton electrons can then deposit energy up to a MIP (2.4 MeV) in each plane of the scintillator, though the amount deposited in each plane will depend on where the Compton scatter occurs.

Lightning Mapping Array. As shown in Figure 1, the LMA consisted of nine stations located within and around the TASD and determines accurate 3-D observations of peak VHF radiation events above threshold in $80 \mu\text{s}$ time intervals (Rison et al., 1999; Thomas et al., 2004). In addition to showing the large scale structure and development of flashes and the lightning flashing rate, its observations were used to determine the plan distance to the TGF events and also to finely calibrate the INTF azimuth and elevation values. The angular calibration was done separately on a flash-by-flash basis for each TGF event.

VHF Lightning Interferometer (INTF) and Fast Electric Field Change Antenna (FA). The INTF records broadband (20–80 MHz) waveforms at 180 MHz from three flat-plate receiving antennas and determines the two-dimensional azimuth and elevation arrival directions of the VHF radiation with sub-microsecond resolution (Stock et al., 2014). This is done on a post-processed basis and determines the radiation centroid in overlapping 0.7 or $1.4 \mu\text{s}$ windows. Triangular baselines of 106–121 m were used to maximize the angular resolution over the TASD. The elevation angles were used to determine the source altitude of the TGFs, based on the LMA-determined plan distance to the source, and the amplitude of the

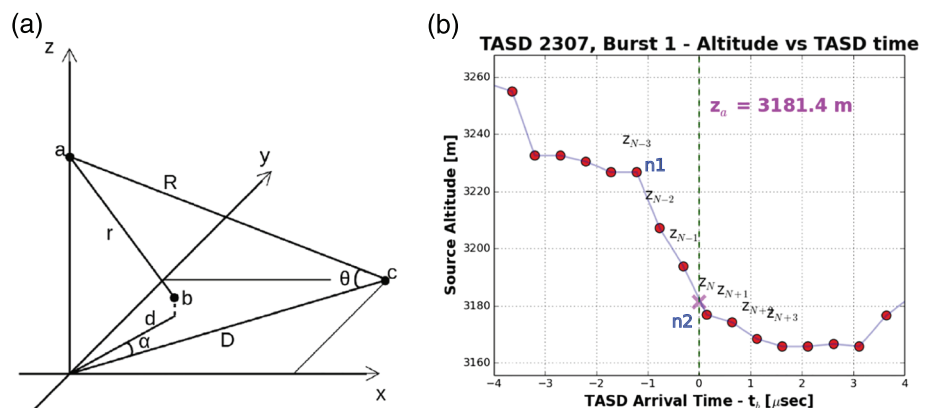


Figure A1. Methods information. (a) Source-centric coordinate system for temporal correlations. The TGF source is at (x_a, y_a, z_a) , with the plan x, y location serving as the coordinate origin. The TASD station is at location b relative to the origin and the reference altitude, and the INTF/FA is at the more distant location c . (b) Iteration at $0.5 \mu\text{s}$ time steps used in the alternative approach for determining the source altitude (TGF A in this case), showing the occurrence of enhanced-speed downward FNB immediately before the TGF onset (red “x”).

received signals was used to determine the VHF power of the centroids. The fast electric field change antenna (FA) provided high resolution (180 MHz) measurements of the low frequency (LF/ELF) discharge sferics that are key to interpreting the INTF and LMA observations.

A2. Analysis Procedures

Figure A1a shows the coordinate system used for analyzing the INTF and T ASD observations. For simplicity, this is done in a Cartesian coordinate system centered at the x_a, y_a plan location of the TGF's source. The plan location is determined from the mean values of the latitude and longitude of LMA sources within ± 1 ms of the TGF's occurrence, seen in Figures 10e–13e. The altitude values are determined relative to a 1,400 m reference plane, which is within 2 m of the GPS altitude of the VHF receiving antenna used as the INTF's GPS time base. The plan locations and altitudes of the T ASD stations are precisely known and fully accounted for in the calculations, with trigger times of each T ASD's data accurate to 40 ns. Similarly, the INTF source directions were carefully calibrated to within 0.08° in azimuth and 0.26° in elevation, obtained by comparing accurately located LMA sources with corresponding INTF source directions separately for each flash.

Given the LMA-estimated values of x_a and y_a , two additional measurements are needed to determine the TGF's onset altitude z_a and time t_a . The source altitude can be estimated from the LMA observations but has insufficient accuracy and temporal resolution to resolve the fast downward breakdown that occurs during the parent IBP (typically 100–150 m in 5–10 μ s). Instead, the altitude is more accurately determined from the INTF elevation angle θ_c versus time, which is obtained with sub-microsecond resolution. In particular, $z_a = D \tan \theta(t_c) = z_a(t_c)$, where $D = \sqrt{x_c^2 + y_c^2}$ is the plan distance between the INTF and TGF. For an event at altitude z_a and time t_a , the arrival times at T ASD i and the INTF are given by

$$t_b = t_a + r_b/c, \quad (\text{A1})$$

$$t_c = t_a + R/c, \quad (\text{A2})$$

where $r_b = [x_i^2 + y_i^2 + (z_a - z_i)^2]^{1/2}$ and $R = [x_c^2 + y_c^2 + z_a^2]^{1/2}$ are the slant ranges from the TGF source. Because the plan locations are considered to be known, $r_b = r_b(z_a)$ and $R = R(z_a)$, so the time-of-arrival equations represent two equations and two unknowns, t_a and z_a . The unknowns are determined from two measurements, in particular the arrival time t_b at a given T ASD station and the INTF elevation measurements, $\theta_c(t_c)$. Since θ_c varies with time during the IBPs, it is not known in advance which time value t_c to use for determining z_a . This results from z_a depending on itself in a manner that is not amenable to analytical inversion. But the equations are readily solved by iterating over the range of values for z_a or equivalently over the possible θ_c or t_c values.

Two semi-independent approaches were used to determine the solutions. Both used an alternative form of (A2) obtained by eliminating t_a to obtain

$$t_c = t_b + \frac{(R - r_b)}{c} = t_b + \Delta t_b, \quad (\text{A3})$$

where $\Delta t_b = (R/c) - (r_b/c)$ corresponds to the time shift for comparing a given T ASD's observations with the INTF/FA observations. For an assumed source altitude z_a , the time shift between the onset time t_b at a given T ASD station and its arrival time t_c at the INTF is readily calculated from the difference of the slant ranges R and r_b of the source relative to the INTF and the T ASD in question. In turn, the t_c value can be used to determine $\theta_c(t_c)$ and hence z_a . Comparing the assumed and inferred z_a values forms the basis for a closed loop iteration procedure in which the assumed z_a is simply replaced by the new z_a value (Figure S14). Consistency is reached in just a few steps. At the same time, the corresponding INTF elevation angle θ_c and arrival time t_c at the INTF are also determined.

The above is the method used by the first approach, as described in section 2.2. For each of the primary TGFs shown in Figure 4, the source altitudes inferred from the onset times at the different T ASDs were in good agreement, having uncertainties of 30, 16, 10, and 40 m for TGFs A, B, C, and D, respectively (see Table S2). To guard against outliers, median values were used for determining the final z_a and t_a values at onset, as well as θ_c and t_c . The final t_c values provide a reference time for evaluating the onset times of each gamma-ray event. As can be seen from the T ASD plots in Figure 4, in most cases the waveforms begin

within a microsecond or so of the indicated t_c onset time. Detections that begin in advance of or after the indicated onset, as for TGF B, are indicative of different onset times.

Instead of using a closed-loop iteration process, as above, the second approach worked backward from the INTF observations of the elevation angle θ_c versus t_c to determine z_a and Δt in reverse. This was used to predict the arrival times at two of the T ASD stations that detected the TGF most strongly and involved stepping through the t_c times and corresponding θ_c values in $0.5 \mu\text{s}$ increments and determining the time when the difference between the predicted and observed t_b values passed through zero. The common reference time t_b was defined to be when the T ASD signal first ascended to half of its eventual peak amplitude on the two stations with the strongest signals (short vertical dotted lines in the T ASD waveforms of Figures S10–S13c and S13d), which were averaged to obtain the final estimate of the time alignment.

Figure A1b shows the results of the stepping procedure for T ASD 2307 of TGF A. The plot shows the difference between the observed and trial t_b times of the main gamma-ray event, with the interpolated step value, where Δt_b goes through zero determining the value of t_c (red “x” in the figure). For this (and the iterative) procedure to work, the INTF data were processed with higher time resolution and increased overlap to make $\theta_c(t_c)$ more continuous. This is a standard procedure for analyzing INTF observations (Stock et al., 2014) and allows more detail to be seen in θ_c versus time. For these analyses, the higher resolution data were down-sampled to $0.5 \mu\text{s}$ intervals by using the median of the higher frequency processing over a $\pm 4 \mu\text{s}$ interval around each $0.5 \mu\text{s}$ point (unfilled gray circles in panels c and d of Figures S10–S13).

What is informative and notable about the example of Figure A1b is that the onset time of the strong gamma burst of TGF A coincided with the end of a brief interlude of rapid descent in the source altitude, denoted by the vertical dashed line in the figure. The speed of the descent is determined from the spacing between the dots, which occur at $0.5 \mu\text{s}$ intervals. In $1.5 \mu\text{s}$ (three step intervals), the source descended about 50 m, corresponding to a downward speed of $3.3 \times 10^7 \text{ m/s}$. This enhanced-speed interlude was unresolved by the normal processing and instead caused the step discontinuity seen in Figure 4a during the fast negative breakdown. The stepping method of determining the onset time agreed well with the result of the iterative approach, which showed the gamma-ray onset to be at the end of the discontinuity (bold vertical line in Figure 4a). The agreement is not surprising, given that the same basic data were used in the two analysis approaches. But the correspondence with different approaches indicates good precision in the procedures and reinforces the observation that the gamma bursts occur in association with intervals of enhanced speed breakdown.

A3. Measurement Uncertainties

Whereas the INTF and FA data are well-synchronized timewise by being simultaneously digitized at a high rate, the main question is how accurately the T ASD waveforms from the different T ASDs are synchronized with the INTF/FA data. As discussed above, this can be qualitatively determined by examining the waveforms from the different SDs relative to the inferred onset time (vertical line) for each of the TGF events in Figure 4. In most cases, the observed onsets are within a microsecond or less of the inferred time, with important exceptions in TGFs B and C.

A quantitative result can only be obtained from propagating the measurements' standard errors through calculations in the previous section, using the general form of

$$\delta f = \sqrt{\left(\frac{\partial f}{\partial x_1} \delta x_1\right)^2 + \dots + \left(\frac{\partial f}{\partial x_n} \delta x_n\right)^2}, \quad (\text{A4})$$

where $f = f(x_1, \dots, x_n)$. Detector locations are known to centimeter accuracy and have negligible contributions. Similarly, gamma-ray detection trigger times are known on the order of sampling rate (10s of ns). Both are taken into account but have very little effect on final uncertainties. Primary error sources, then, come from the two instances of taking averages described the previous section; TGF source plan locations are taken as the mean GPS location of LMA sources within 1 ms of particle detections, and its uncertainty is the standard error. TGF source elevations are done the same way—a mean is taken of all INTF sources within $4 \mu\text{s}$ of the TGFs inferred arrival at the interferometer (from equation A3), and its uncertainty is the standard error.

All subsequent calculations can then be shadowed by their error counterparts using equation (A4) and are presented in Tables S2 and S3. Typically, altitude measurements are much less precise for this type of study,

but here altitude determination comes from the higher-sampled INTF data whereas plan location data are supplied by only a few LMA points. As a result, altitude uncertainties are 30, 20, 10, and 40 m for TGFs A, B, C, and D, respectively, compared to horizontal location errors of 150, 80, 40, and 300 m. Timing uncertainties follow the same trend, with 0.7, 0.4, 0.2, and 1.4 μ s for each respective TGF. Standard errors for all other calculations are shown in Tables S2 and S3.

Acknowledgments

The lightning instrumentation, operation, and analyses of this study have been supported by NSF grants AGS-1205727, AGS-1613260, AGS-1720600, and AGS-1844306. The Telescope Array experiment is supported by the Japan Society for the Promotion of Science through Grants-in-Aids for Scientific Research on Specially Promoted Research (15H05693) and for Scientific Research (S) (15H05741) and the Inter-University Research Program of the Institute for Cosmic Ray Research; by the U.S. National Science Foundation awards PHY-0307098, PHY-0601915, PHY-0649681, PHY-0703893, PHY-0758342, PHY-0848320, PHY-1069280, PHY-1069286, PHY-1404495, PHY-1404502, and PHY-1607727; by the National Research Foundation of Korea (2015R1A2A1A01006870, 2015R1A2A1A15055344, 2016R1A5A1013277, 2007-0093860, 2016R1A2B4014967, and 2017K1A4A3015188); by the Russian Academy of Sciences, RFBR grant 16-02-00962a (INR), IISN project No. 4.4502.13, and Belgian Science Policy under IUAP VII/37 (ULB). The foundations of Dr. Ezekiel R. and Edna Wattis Dumke, Willard L. Eccles, and George S. and Dolores Doré Eccles all helped with generous donations. The State of Utah supported the project through its Economic Development Board and the University of Utah through the Office of the Vice President for Research. The experimental site became available through the cooperation of the Utah School and Institutional Trust Lands Administration (SITLA), U.S. Bureau of Land Management (BLM), and the U.S. Air Force. We appreciate the assistance of the State of Utah and Fillmore offices of the BLM in crafting the Plan of Development for the site. We also wish to thank the people and the officials of Millard County, Utah for their steadfast and warm support. We gratefully acknowledge the contributions from the technical staffs of our home institutions. An allocation of computer time from the Center for High Performance Computing at the University of Utah is gratefully acknowledged. We thank Ryan Said and W. A. Brooks of Vaisala Inc. for providing high-quality NLDN data lightning discharges over and around the TASD under their academic research use policy as well as several anonymous reviewers for their requests and comments, the responses to which significantly improved the paper.

Notice that elevation errors are nearly equal (Table S2), but poor grouping of LMA data at the time of TGF D means a larger error in the plan location. As the error is propagated through each calculation, quantities for TGF D continue to be the least reliable among the four, showing that the low LMA sampling rate and possible mislocations during fast breakdown are the main contributors to all further uncertainty.

Data Availability Statement

Data that support the conclusions presented in the manuscript are provided in the figures of the paper. Additional information can be found in the supporting material and is available on the Open Science Framework (<https://doi.org/10.17605/OSF.IO/Z3XDA>).

References

Abbasi, R., Abe, M., Abu-Zayyad, T., Allen, M., Anderson, R., Azuma, R., et al. (2017). The bursts of high energy events observed by the telescope array surface detector. *Physics Letters A*, 381(32), 2565–2572. <https://doi.org/10.1016/j.physleta.2017.06.022>

Abbasi, R. U., Abu-Zayyad, T., Allen, M., Barcikowski, E., Belz, J. W., Bergman, D. R., et al. (2018). Gamma ray showers observed at ground level in coincidence with downward lightning leaders. *Journal of Geophysical Research: Atmospheres*, 123, 6864–6879. <https://doi.org/10.1029/2017JD027931>

Abu-Zayyad, T., Aida, R., Allen, M., Anderson, R., Azuma, R., Barcikowski, E., et al. (2013). The surface detector array of the Telescope Array experiment. *Nuclear Instrumentation and Methods in Physics Research*, A689, 87–97. <https://doi.org/10.1016/j.nima.2012.05.079>

Attanasio, A., Krehbiel, P., & da Silva, C. (2019). Griffiths and Phelps lightning initiation model, revisited. *Journal of Geophysical Research: Atmospheres*, 124, 8076–8094. <https://doi.org/10.1029/2019JD030399>

Behnke, S. A., Thomas, R. J., Krehbiel, P. R., & Rison, W. (2005). Initial leader velocities during intracloud lightning: Possible evidence for a runaway breakdown effect. *Journal of Geophysical Research*, 110, D10207. <https://doi.org/10.1029/2004JD005312>

Biagi, C. J., Uman, M. A., Hill, J. D., Jordan, D. M., Rakov, V. A., & Dwyer, J. (2010). Observations of stepping mechanisms in a rocket-and-wire triggered lightning flash. *Journal of Geophysical Research*, 115, D23215. <https://doi.org/10.1029/2010JD014616>

Briggs, M. S., Fishman, G. J., Connaughton, V., Bhat, P. N., Paciesas, W. S., Preece, R. D., et al. (2010). First results on terrestrial gamma ray flashes from the Fermi gamma-ray burst monitor. *Journal of Geophysical Research*, 115, A07323. <https://doi.org/10.1029/2009JA015242>

Briggs, M. S., Xiong, S., Connaughton, V., Tierney, D., Fitzpatrick, G., Foley, S., et al. (2013). Terrestrial gamma-ray flashes in the Fermi era: Improved observations and analysis methods. *Journal of Geophysical Research: Space Physics*, 118, 3805–3830. <https://doi.org/10.1002/jgra.50205>

Carlson, B. E., Lehtinen, N. G., & Inan, U. S. (2007). Constraints on terrestrial gamma ray flash production from satellite observation. *Geophysical Research Letters*, 34, L08809. <https://doi.org/10.1029/2006GL029229>

Celestin, S., & Pasko, V. (2011). Energy and fluxes of thermal runaway electrons produced by exponential growth of streamers during the stepping of lightning leaders and in transient luminous events. *Journal of Geophysical Research*, 116, A03315. <https://doi.org/10.1029/2010JA016260>

Celestin, S., & Pasko, V. (2012). Compton scattering effects on the duration of terrestrial gamma-ray flashes. *Geophysical Research Letters*, 39, L02802. <https://doi.org/10.1029/2011GL050342>

Celestin, S., Xu, W., & Pasko, V. (2012). Terrestrial gamma ray flashes with energies up to 100 MeV produced by nonequilibrium acceleration of electrons in lightning. *Journal of Geophysical Research: Space Physics*, 117. <https://doi.org/10.1029/2012JA01753>

Celestin, S., Xu, W., & Pasko, V. (2015). Variability in fluence and spectrum of high-energy photon bursts produced by lightning leaders. *Journal of Geophysical Research: Space Physics*, 120, 10,712–10,723. <https://doi.org/10.1002/2015JA021410>

Cummer, S. A. (2020). Indirectly measured ambient electric fields for lightning initiation in fast breakdown regions. *Geophysical Research Letters*, 47, e2019GL086089. <https://doi.org/10.1029/2019GL086089>

Cummer, S. A., Lu, G., Briggs, M. S., Connaughton, V., Xiong, S., Fishman, G. J., & Dwyer, J. R. (2011). The lightning-TGF relationship on microsecond timescales. *Geophysical Research Letters*, 38, L14810. <https://doi.org/10.1029/2011GL048099>

Cummer, S. A., Lyu, F., Briggs, M. S., Cramer, E., Stanbro, M., Roberts, O., & Smith, D. M. (2017). *The connection between terrestrial gamma-ray flashes and energetic in-cloud lightning pulses*, (Abstract AE33B-2547). Paper presented at American Geophysical Union Fall Meeting, New Orleans.

Cummer, S. A., Lyu, F., Briggs, M. S., Fitzpatrick, G., Roberts, O. J., & Dwyer, J. R. (2015). Lightning leader altitude progression in terrestrial gamma-ray flashes. *Geophysical Research Letters*, 42, 7792–7798. <https://doi.org/10.1002/2015GL065228>

da Silva, C., & Pasko, V. (2013). Dynamics of streamer-to-leader transition at reduced air densities and its implications for propagation of lightning leaders and gigantic jets. *Journal of Geophysical Research: Atmospheres*, 118, 13,561–13,590. <https://doi.org/10.1002/2013JD020618>

da Silva, C., & Pasko, V. (2015). Physical mechanism of initial breakdown pulses and narrow bipolar events in lightning discharges. *Journal of Geophysical Research: Atmospheres*, 120, 4989–5009. <https://doi.org/10.1002/2015JD023209>

da Silva, C., Sonnenfeld, R. G., Edens, H. E., Krehbiel, P. R., Quick, M. G., & Koshak, W. J. (2019). The plasma nature of lightning channels and the resulting nonlinear resistance. *Journal of Geophysical Research: Atmospheres*, 124, 9442–9463. <https://doi.org/10.1029/2019JD030693>

Dwyer, J. (2003). A fundamental limit on electric fields in air. *Geophysical Research Letters*, 30(20), 2055. <https://doi.org/10.1029/2003GL017781>

Dwyer, J. (2004). Implications of x-ray emission from lightning. *Geophysical Research Letters*, 31, L12102. <https://doi.org/10.1029/2004GL019795>

- Dwyer, J. (2012). The relativistic feedback discharge model of terrestrial gamma ray flashes. *Journal of Geophysical Research*, *117*, A02308. <https://doi.org/10.1029/2011JA017160>
- Dwyer, J., Schaal, M. M., Cramer, E., Arabshahi, S., Liu, N., Rassoul, H. K., et al. (2012). Observation of a gamma-ray flash at ground level in association with a cloud-to-ground lightning return stroke. *Journal of Geophysical Research*, *117*, A10303. <https://doi.org/10.1029/2012JA017810>
- Edens, H. E., Eack, K. B., Rison, W., & Hunyady, S. J. (2014). Photographic observations of streamers and steps in a cloud-to-air negative leader. *Geophysical Research Letters*, *41*(4), 1336–1342. <https://doi.org/10.1002/2013GL059180>
- Fishman, G. J., Bhat, P. N., Mallozzi, R., Horack, J. M., Koshut, T., Kouveliotou, C., et al. (1994). Discovery of intense gamma-ray flashes of atmospheric origin. *Science*, *264*(5163), 1313–1316. <https://doi.org/10.1126/science.264.5163.1313>
- Foley, S., Fitzpatrick, G., Briggs, M. S., Connaughton, V., Tierney, D., McBreen, S., et al. (2014). Pulse properties of terrestrial gamma-ray flashes detected by the Fermi Gamma-Ray Burst Monitor. *Journal of Geophysical Research: Space Physics*, *119*, 5931–5942. <https://doi.org/10.1002/2014JA019805>
- Gjesteland, T., Østgaard, N., Collier, A. B., Carlson, B. E., Eyles, C., & Smith, D. M. (2012). A new method reveals more TGFs in the RHESSI data. *Geophysical Research Letters*, *39*, L05102. <https://doi.org/10.1029/2012GL050899>
- Grefenstette, B. W., Smith, D. M., Hazelton, B. J., & Lopez, L. I. (2009). First RHESSI terrestrial gamma ray flash catalog. *Journal of Geophysical Research*, *114*, A02314. <https://doi.org/10.1029/2008JA013721>
- Hare, B. M., Uman, M. A., Dwyer, J. R., Jordan, D. M., Biggstaff, M. I., Caicedo, J. A., et al. (2016). Ground-level observation of a terrestrial gamma-ray flash initiated by triggered lightning. *Journal of Geophysical Research: Atmospheres*, *121*, 6511. <https://doi.org/10.1002/2015JD024426>
- Karunarathne, N., Karunarathne, S., Marshall, T. C., & Stolzenburg, M. (2014). Modeling initial breakdown pulses of CG lightning flashes. *Journal of Geophysical Research: Atmospheres*, *119*, 9003. <https://doi.org/10.1002/2014JD021553>
- Krehbiel, P. (2018). *The initial development of intracloud flashes* (Abstract AE11A-01). Paper presented at American Geophysical Union Fall Meeting, Washington, D.C.
- Krehbiel, P. R., Rioussel, J. A., Pasko, V. P., Thomas, R. J., Rison, W., Stanley, M. A., & Edens, H. E. (2008). Upward electrical discharges from thunderstorms. *Nature Geoscience*, *1*, 233–237. <https://doi.org/10.1038/ngeo162>
- LeVine, D. (1980). Sources of the strongest RF radiation from lightning. *Journal of Geophysical Research*, *85*(C7), 4091–4095. <https://doi.org/10.1029/JC085iC07p04091>
- Lu, G., Blakeslee, R. J., Li, J., Smith, D. M., Shao, X.-M., McCaul, E. W., et al. (2010). Lightning mapping observation of a terrestrial gamma ray flash. *Geophysical Research Letters*, *37*, L11806. <https://doi.org/10.1029/2010GL043494>
- Lu, G., Cummer, S. A., Li, J., Han, F., Smith, D. M., & Grefenstette, B. W. (2011). Characteristics of broadband lightning emissions associated with terrestrial gamma ray flashes. *Journal of Geophysical Research*, *116*, A03316. <https://doi.org/10.1029/2010JA016141>
- Lyu, F., Cummer, S. A., Briggs, M., Marisaldi, M., Blakeslee, R. J., Bruning, E., et al. (2016). Ground detection of terrestrial gamma ray flashes from distant radio signals. *Geophysical Research Letters*, *43*, 8728–8734. <https://doi.org/10.1002/2016GL070154>
- Lyu, F., Cummer, S. A., Krehbiel, P. R., Rison, W., Briggs, M. S., Cramer, E., et al. (2018). Very high frequency radio emissions associated with the production of terrestrial gamma-ray flashes. *Geophysical Research Letters*, *45*, 2097–2105. <https://doi.org/10.1002/2018GL077102>
- Lyu, F., Cummer, S. A., & McTague, L. (2015). Insights into high peak current in-cloud lightning events during thunderstorms. *Geophysical Research Letters*, *42*, 6836–6843. <https://doi.org/10.1002/2015GL065047>
- MacGorman, D. R., & Rust, W. D. (1998). *The electrical nature of storms*. New York: Oxford University Press. Retrieved from https://books.google.com/books?id=_NbHNj7KJecC
- Mailyan, B. G., Briggs, M. S., Cramer, E. S., Fitzpatrick, G., Roberts, O. J., Stanbro, M., et al. (2016). The spectroscopy of individual terrestrial gamma-ray flashes: Constraining the source properties. *Journal of Geophysical Research: Space Physics*, *121*, 11,346–11,363. <https://doi.org/10.1002/2016JA022702>
- Mailyan, B. G., Nag, A., Murphy, M. J., Briggs, M. S., Dwyer, J. R., Rison, W., et al. (2018). Characteristics of radio emissions associated with terrestrial gamma-ray flashes. *Journal of Geophysical Research: Space Physics*, *123*, 5933–5948. <https://doi.org/10.1029/2018JA025450>
- Mailyan, B. G., Xu, W., Celestin, S., Briggs, M. S., Dwyer, J. R., Cramer, E. S., et al. (2019). Analysis of individual terrestrial gamma-ray flashes with lightning leader models and fermi gamma-ray burst monitor data. *Journal of Geophysical Research: Atmospheres*, *124*, 7170–7183. <https://doi.org/10.1029/2019JA026912>
- Malagon-Romero, A., & Luque, A. (2019). Spontaneous emergence of space stems ahead of negative leaders in lightning and long sparks. *Geophysical Research Letters*, *46*, 4029–4038. <https://doi.org/10.1029/2019GL082063>
- Marisaldi, M., Fuschino, F., Tavani, M., Dietrich, S., Price, C., Galli, M., et al. (2014). Properties of terrestrial gamma ray flashes detected by AGILE MCAL below 30 MeV. *Journal of Geophysical Research: Space Physics*, *119*, 1337–1355. <https://doi.org/10.1002/2013JA019301>
- Marshall, T., Stolzenburg, M., Karunarathne, S., Cummer, S., Lu, G., Betz, H.-D., et al. (2013). Initial breakdown pulses in intracloud lightning flashes and their relation to terrestrial gamma ray flashes. *Journal of Geophysical Research: Atmospheres*, *118*, 10,907–10,925. <https://doi.org/10.1002/jgrd.50866>
- Moss, G., Pasko, V., Liu, N., & Veronis, G. (2006). Monte carlo model for analysis of thermal runaway electrons in streamer tips in transient luminous events and streamer zones of lightning leaders. *Journal of Geophysical Research*, *111*, A02307. <https://doi.org/10.1029/2005JA011350>
- Nag, A., DeCarlo, B., & Rakov, V. (2009). Analysis of microsecond- and submicrosecond-scale electric field pulses produced by cloud and ground lightning discharges. *Atmospheric Research*, *91*, 316. <https://doi.org/10.1016/j.atmosres.2008.01.014>
- Nijdam, S., Teunissen, J., & Ebert, U. (2020). The physics of streamer discharge phenomena. *Plasma Sources Science and Technology*. Retrieved from <https://iopscience.iop.org/article/10.1088/1361-6595/abaa05>
- Østgaard, N., Neubert, T., Reglero, V., Ullaland, K., Yang, S., Genov, G., et al. (2019). First 10 months of TGF observations by ASIM. *Journal of Geophysical Research: Atmospheres*, *124*, 14,024–14,036. <https://doi.org/10.1029/2019JD031214>
- Petersen, D., Bailey, M., Beasley, W., & Hallett, J. (2008). A brief review of the problem of lightning initiation and a hypothesis of initial lightning leader formation. *Journal of Geophysical Research*, *113*, D17205. <https://doi.org/10.1029/2007JD009036>
- Pu, Y., Cummer, S. A., Lyu, F., Briggs, M., Mailyan, B., Stanbro, M., & Roberts, O. (2019). Low frequency radio pulses produced by terrestrial gamma ray flashes. *Geophysical Research Letters*, *46*, 6990–6997. <https://doi.org/10.1029/2019GL082743>
- Ringuette, R., Case, G. L., Cherry, M. L., Granger, D., Guzik, T. G., Stewart, M., & Wefel, J. P. (2013). TETRA observation of gamma-rays at ground level associated with nearby thunderstorms. *Journal of Geophysical Research: Space Physics*, *118*, 7841. <https://doi.org/10.1002/jgra.50712>
- Rison, W., Krehbiel, P., Stock, M., Edens, H., Shao, X., Thomas, R., et al. (2016). Observations of narrow bipolar events reveal how lightning is initiated in thunderstorms. *Nature Communications*, *7*, 10721. <https://doi.org/10.1038/ncomms10721>

- Rison, W., Thomas, R., Krehbiel, P., Hamlin, T., & Harlin, J. (1999). A GPS-based three-dimensional lightning mapping system: Initial observations in central New Mexico. *Geophysical Research Letters*, *26*(23), 3573–3576. <https://doi.org/10.1029/1999GL010856>
- Roberts, O. J., Fitzpatrick, G., Priftis, G., Bedka, K., Chronis, T., McBreen, S., et al. (2017). Terrestrial gamma ray flashes due to particle acceleration in tropical storm systems. *Journal of Geophysical Research: Atmospheres*, *122*, 3374–3395. <https://doi.org/10.1002/2016JD025799>
- Shao, X., Hamlin, T., & Smith, D. (2010). A closer examination of terrestrial gamma ray flash related lightning processes. *Journal of Geophysical Research: Space Physics*, *115*, A00E30. <https://doi.org/10.1029/2009JA014835>
- Shi, F., Liu, N., & Rassoul, H. K. (2016). Properties of relatively long streamers initiated from an isolated hydrometeor. *Journal of Geophysical Research: Atmospheres*, *121*, 7284–7295. <https://doi.org/10.1002/2015JD024580>
- Shi, D., Wang, D., Wu, T., & Takagi, N. (2019). Temporal and spatial characteristics of preliminary breakdown pulses in intracloud lightning flashes. *Journal of Geophysical Research: Atmospheres*, *124*, 12,901–12,914. <https://doi.org/10.1029/2019JD031130>
- Smith, D. M., Bowers, G. S., Kamogawa, M., Wang, D., Ushio, T., Ortberg, J., et al. (2018). Characterizing upward lightning with and without a terrestrial gamma ray flash. *Journal of Geophysical Research: Atmospheres*, *123*, 11,321–11,332. <https://doi.org/10.1029/2018JD029105>
- Stanley, M. A., Shao, X.-M., Smith, D. M., Lopez, L. I., Pongratz, M. B., Harlin, J. D., et al. (2006). A link between terrestrial gamma-ray flashes and intracloud lightning discharges. *Geophysical Research Letters*, *33*, L06803. <https://doi.org/10.1029/2005GL025537>
- Stock, M. G., Akita, M., Krehbiel, P. R., Rison, W., Edens, H. E., Kawasaki, Z., & Stanley, M. A. (2014). Continuous broadband digital interferometry of lightning using a generalized cross-correlation algorithm. *Journal of Geophysical Research: Atmospheres*, *119*, 3134–3165. <https://doi.org/10.1002/2013JD020217>
- Stolzenburg, M., Marshall, T. C., Karunarathne, S., Karunarathna, N., Vickers, L. E., Warner, T. A., et al. (2013). Luminosity of initial breakdown in lightning. *Journal of Geophysical Research*, *118*, 2918–2937. <https://doi.org/10.1002/jgrd.50276>
- Stolzenburg, M., Marshall, T. C., Karunarathne, S., Karunarathna, N., & Orville, R. E. (2014). Leader observations during the initial breakdown stage of a lightning flash. *Journal of Geophysical Research: Atmospheres*, *119*, 12,198–12,221. <https://doi.org/10.1002/2014JD021994>
- Stolzenburg, M., Marshall, T. C., Karunarathne, S., & Orville, R. E. (2016). Luminosity with intracloud-type lightning initial breakdown pulses and terrestrial gamma-ray flash candidates. *Journal of Geophysical Research: Atmospheres*, *121*, 10,919–10,936. <https://doi.org/10.1002/2016JD025202>
- Thomas, R. J., Krehbiel, P. R., Rison, W., Hunyady, S. J., Winn, W. P., Hamlin, T., & Harlin, J. (2004). Accuracy of the Lightning Mapping Array. *Journal of Geophysical Research: Atmospheres*, *109*, D14207. <https://doi.org/10.1029/2004JD004549>
- Tilles, J. (2020). Broadband radio mapping and imaging of lightning processes (PhD thesis), University of New Hampshire.
- Tilles, J. N., Krehbiel, P. R., Stanley, M. A., Rison, W., Liu, N., Lyu, F., et al. (2020). Radio interferometer observations of an energetic in-cloud pulse reveal large currents generated by relativistic discharges. *Journal of Geophysical Research: Atmospheres*, *125*, e2020JD032603. <https://doi.org/10.1029/2020JD032603>
- Tilles, J. N., Liu, N., Stanley, M. A., Krehbiel, P. R., Rison, W., Stock, M. G., et al. (2019). Fast negative breakdown in thunderstorms. *Nature Communications*, *10*, 1648. <https://doi.org/10.1038/s41467-019-09621-z>
- Tran, M. D., Rakov, V. A., Mallick, S., Dwyer, J. R., Nag, A., & Heckman, S. (2015). A terrestrial gamma-ray flash recorded at the lightning observatory in Gainesville, Florida. *Journal of Atmospheric and Solar-Terrestrial Physics*, *136*, 86. <https://doi.org/10.1016/j.jastp.2015.10.010>
- Wada, Y., Enoto, T., Nakamura, Y., Furuta, Y., Yuasa, T., Nakazawa, K., et al. (2019). Gamma-ray glow preceding downward terrestrial gamma-ray flash. *Physics Communications*, *2*, 67. <https://doi.org/10.1038/s42005-019-0168-y>
- Weidman, C. D., & Krider, E. P. (1979). The radiation field waveforms produced by intracloud lightning discharge processes. *Journal of Geophysical Research*, *84*, 3159. <https://doi.org/10.1029/JC084iC06p03159>
- Winn, W. P., Aulich, G. D., Hunyady, S. J., Eack, K. B., Edens, H. E., Krehbiel, P. R., et al. (2011). Lightning leader stepping, K changes, and other observations near an intracloud flash. *Journal of Geophysical Research*, *116*, D23115. <https://doi.org/10.1029/2011JD015998>
- Zyla, P. A. (2020). Particle data group. *Progress of Theoretical and Experimental Physics*. Retrieved from <https://pdg.lbl.gov/2020/reviews/rpp2020-rev-passage-particles-matter.pdf>